

Climate and vegetation water use efficiency at catchment scales

Peter A. Troch

Surface Water Hydrology
Dept. of Hydrology and Water Resources
The University of Arizona
Tucson, AZ USA

Email: patroch@hwr.arizona.edu
Url: www.hwr.arizona.edu/~surface

Collaborators

- Guillermo Martinez – The University of Arizona
- Valentijn Pauwels – University of Ghent
- Ciaran Harman – University of Illinois at UC
- Murugesu Sivapalan – University of Illinois at UC
- Praveen Kumar – University of Illinois at UC
- Travis Huxman – The University of Arizona
- Hoshin Gupta – The University of Arizona

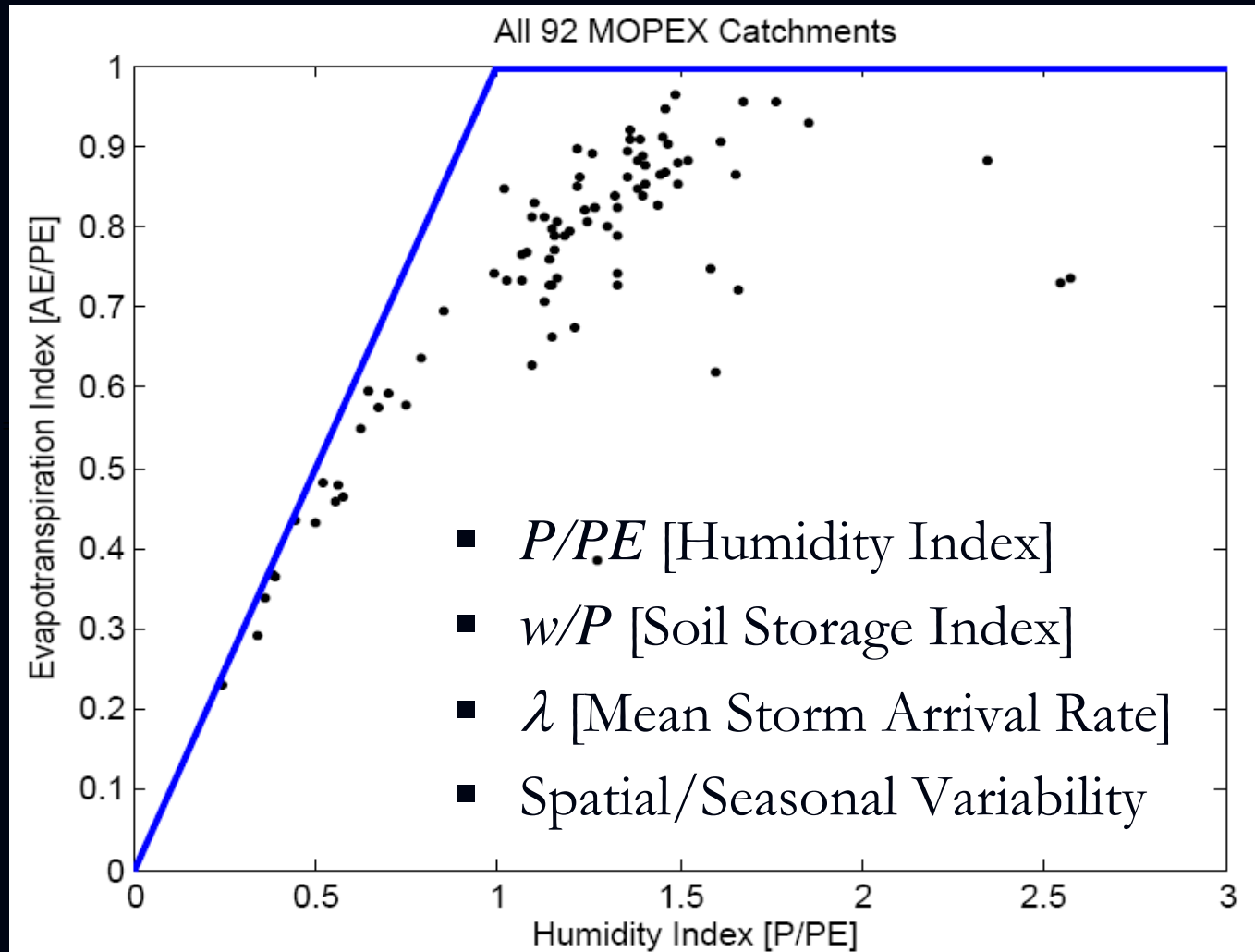
Support

- NSF EAR-Hydrologic Sciences: Understanding the hydrologic implications of landscape structure and climate - Towards a unifying framework of watershed similarity (PIs: Thorsten Wagener, Murugesu Sivapalan, Peter Troch);
- NSF EAR-Hydrologic Sciences: Water cycle dynamics in a changing environment: Advancing hydrologic science through synthesis (PIs: Murugesu Sivapalan, Praveen Kumar, Bruce Roads, Don Wuebbles)

Outline

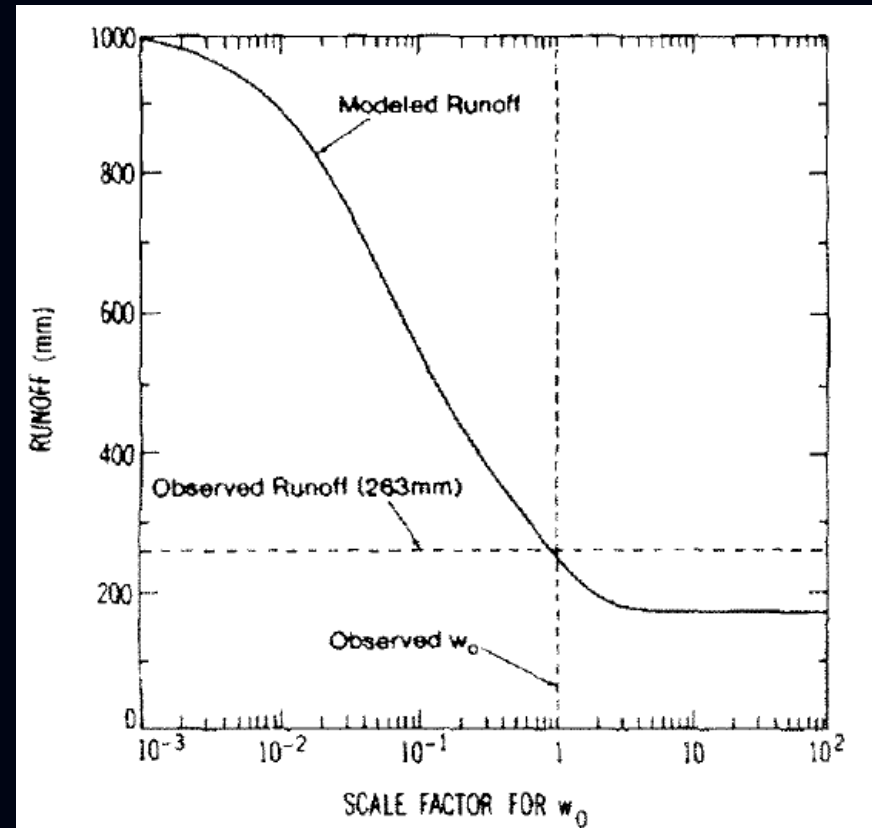
- Background and motivation
- Testing the Horton index
- Precipitation and vegetation productivity
- The annual water balance and L'vovich proportionality relations
- Testing the Ponce and Shetty model
- Conclusions

Budyko's hypothesis: $\frac{V}{PE} = \phi \left(\frac{P}{PE} \right)$



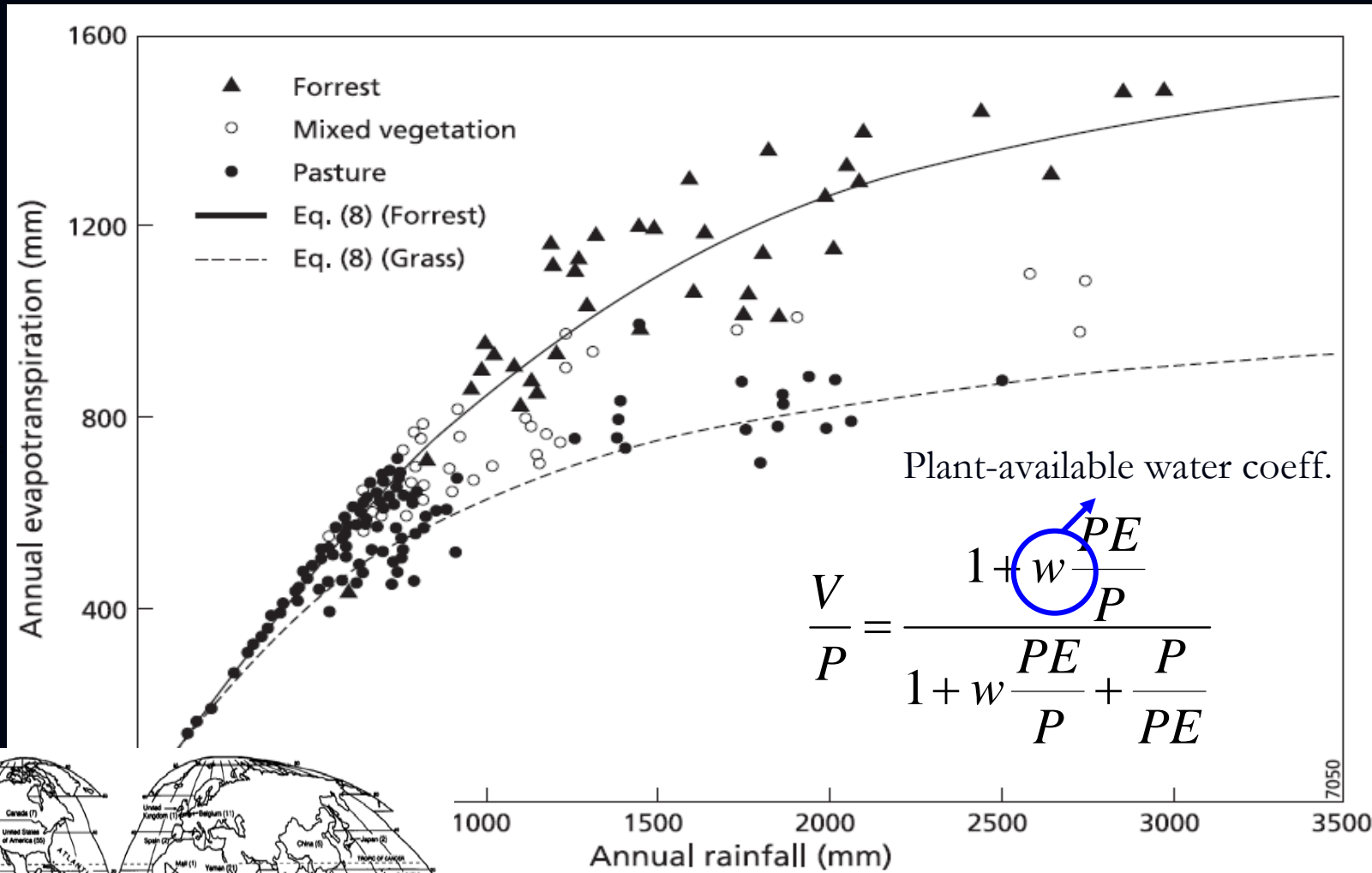
Sensitivity of water balance to water holding capacity

- Sensitivity diminishes at a scale factor on the order of 1;
- This implies that the actual values of capacity are almost large enough to maximize evapo-transpiration (minimize runoff);
- This could indicate that “the rooting depth of plants reflects ecologically optimized responses to the relative timing and magnitude of water and energy supplies”.



Plants are in control?

WUE =



Zhang et al., 2001 (WRR)

Motivation: another Horton index...

REPORTS AND PAPERS, HYDROLOGY--1933

TABLE 3 - SEASONAL RAINFALL, RUNOFF, AND WATER-LOSSES
WEST BRANCH OF DELAWARE RIVER AT HANCOCK AND MALE EDDY, NY
SUMMER SEASON - MAY TO OCTOBER INCL.
(All quantities except ratios are in inches Depth)

| Year | P | Y | Y _g | Y _s | E _p | L' | P-Y | T _{inf} | (%) |
|------|-------|------|----------------|----------------|----------------|-------|-------|------------------|------|
| (1) | (2) | (3) | (4) | (5) | (6) | (7) | (8) | (9) | (10) |
| 1906 | 23.29 | 6.06 | 2.97 | 3.09 | 0.932 | 17.23 | 20.20 | 0.853 | 0.85 |
| 06 | 23.08 | 7.11 | 4.21 | 2.90 | 1.01 | 15.97 | 20.8 | .751 | .79 |
| 07 | 20.30 | 4.89 | 3.55 | 2.85 | 0.928 | 13.47 | 17.45 | .772 | .80 |
| 08 | 16.71 | 5.32 | 2.95 | 2.37 | 1.17 | 11.39 | 14.34 | .784 | .82 |
| 08 | 15.59 | 4.78 | 3.36 | 2.42 | 1.05 | 8.61 | 12.7 | .724 | .74 |
| 1910 | 15.62 | 4.54 | 2.90 | 1.66 | 1.13 | 11.10 | 14.90 | .793 | .80 |
| 11 | 20.65 | 5.42 | 3.09 | 2.33 | 1.06 | 15.23 | 16.72 | .831 | .85 |
| 12 | 18.71 | 4.43 | 3.10 | 1.33 | 1.04 | 14.28 | 17.38 | .822 | .84 |
| 13 | 19.20 | 3.40 | 1.49 | 0.94 | 1.08 | 13.80 | 15.24 | .904 | .93 |
| 14 | 20.90 | 3.47 | 2.55 | 2.12 | 1.04 | 15.23 | 18.78 | .811 | .86 |
| 1915 | 24.90 | 3.89 | 5.12 | 3.45 | 0.947 | 19.01 | 21.65 | .744 | .77 |
| 16 | 29.19 | 3.87 | 5.28 | 2.9 | 1.08 | 15.32 | 24.19 | .715 | .72 |
| 17 | 32.64 | 3.41 | 7.44 | 5.95 | 0.953 | 18.73 | 26.19 | .715 | .72 |
| 18 | 26.68 | 3.62 | 2.40 | 2.22 | 1.08 | 15.98 | 24.38 | .758 | .82 |
| 19 | 21.18 | 3.25 | 4.00 | 1.25 | 1.00 | 15.53 | 19.93 | .799 | .80 |
| 1920 | 27.40 | 3.89 | 5.94 | 2.95 | 1.00 | 18.51 | 24.45 | .757 | .79 |
| 21 | 23.18 | 4.69 | 3.68 | 1.01 | 1.11 | 18.49 | 22.17 | .834 | .83 |
| 22 | 26.09 | 3.28 | 5.35 | 2.93 | 1.01 | 17.75 | 23.10 | .798 | .80 |
| 23 | 21.85 | 3.08 | 3.17 | 1.91 | 1.00 | 14.17 | 15.94 | .841 | .86 |
| 24 | 24.53 | 0.40 | 5.46 | 4.94 | 0.533 | 14.13 | 19.59 | .721 | .82 |
| 25 | 23.02 | 6.14 | 4.62 | 3.32 | 0.882 | 14.68 | 14.70 | .755 | .82 |
| 26 | 20.37 | 5.88 | 3.78 | 2.10 | 0.960 | 14.49 | 14.27 | .793 | .80 |
| 27 | 26.69 | 1.54 | 5.39 | 6.17 | 0.812 | 15.13 | 23.52 | .737 | .82 |
| 28 | 25.35 | 3.06 | 7.54 | 5.51 | 0.880 | 12.25 | 13.64 | .819 | .83 |
| 1929 | 22.39 | 7.18 | 4.98 | 3.22 | 1.01 | 15.21 | 23.17 | .754 | .82 |
| MEAN | 22.46 | 7.27 | 4.42 | 2.85 | | 15.20 | 20.63 | .8177 | 0.90 |

(a) This is apparent infiltration-capacity in inches per day as deduced from daily rainfall-records. The actual infiltration-capacity is greater in the ratio of 24 to the number of hours per day of rainfall-excess duration.

$$H = \frac{V}{W} \cong \text{constant}$$

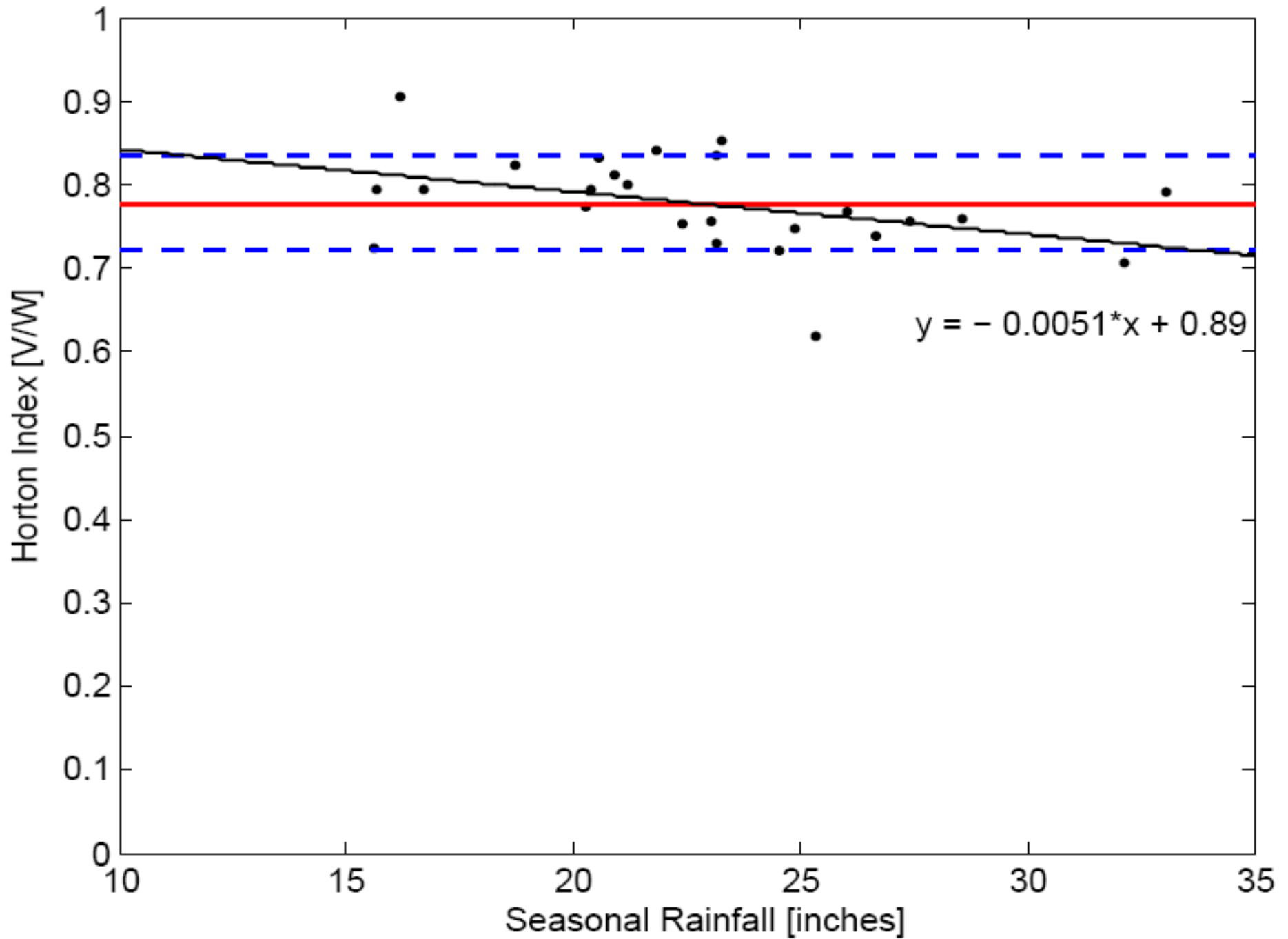
V : Growing-season vaporization (E+T)

W : Growing-season wetting (P-S)

“The natural vegetation of a region tends to develop to such an extent that it can utilize the largest possible proportion of the available soil moisture supplied by infiltration”
(Horton, 1933, p.455)

Horton, 1933 (AGU)

West Branch Delaware River at Hancock, N.Y.



A closer look at the Horton index

$$H = \frac{V}{W} \cong \frac{P - R}{P - S}$$

P : Growing-season rainfall

R : Growing-season total runoff (discharge)

S : Growing-season surface runoff (quick runoff)

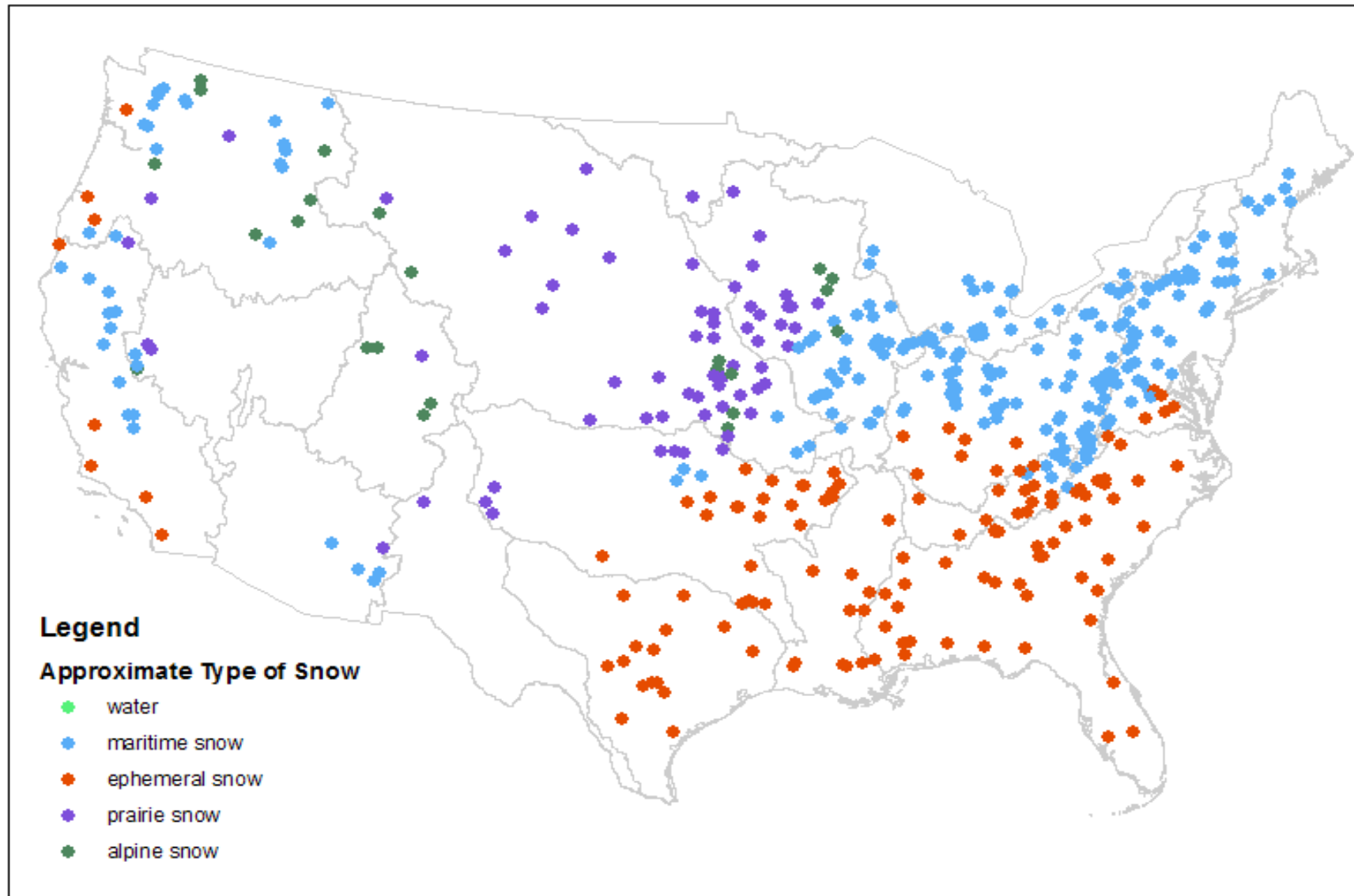
No energy: $P - R = V = 0 : H = 0$

No storage: $R = S = P : H = 0/0$

Humid: $R > S : H < 1$

Semi-arid: $R \cong S < P : H \cong 1$

MOPEX watershed to test Horton Hypothesis

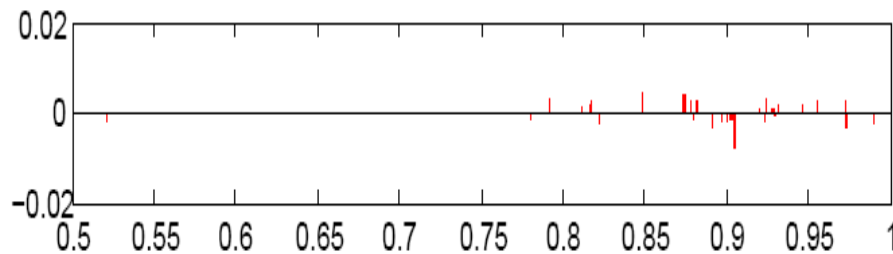
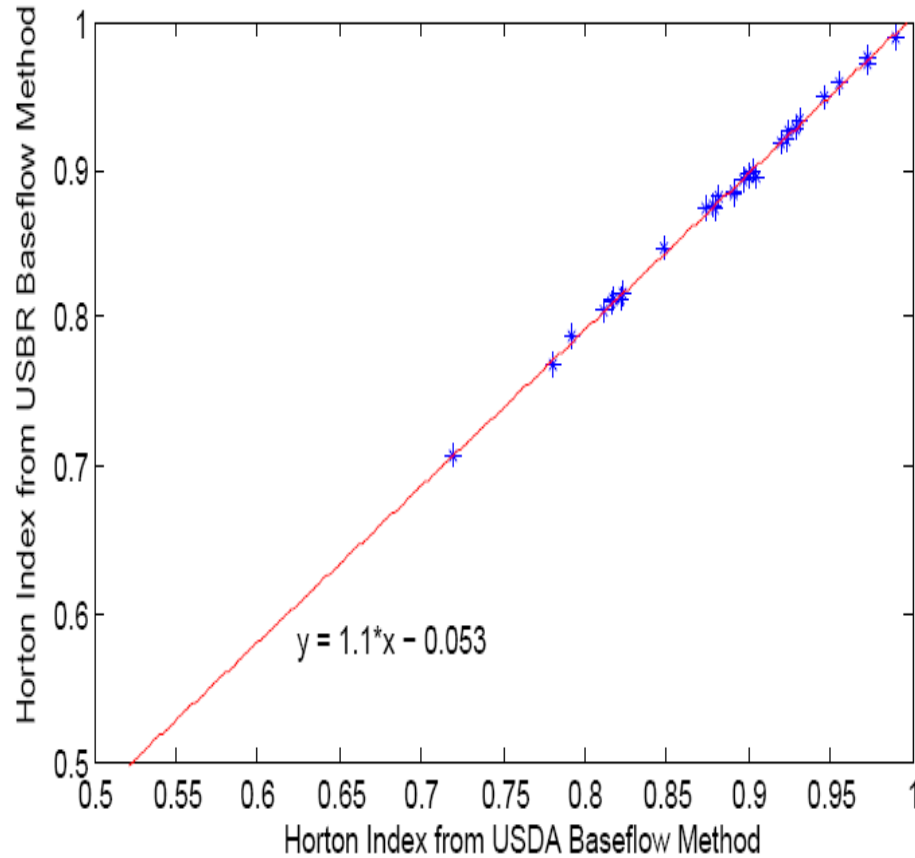


Three Baseflow Separation Methods

- USBR Method (*Wahl and Wahl, 2006*)
 - Based on IH method (recession slope test)
- USDA Method (*Arnold and Allen, 1999*)
 - Method adopted in SWAT model
- UG Method (*Huyck et al., 2005*)
 - Based on hydraulic groundwater theory
 - Accounts for catchment's geomorphology

Comparison of Results

USDA versus USBR Method



USDA versus Huyck et al. Method

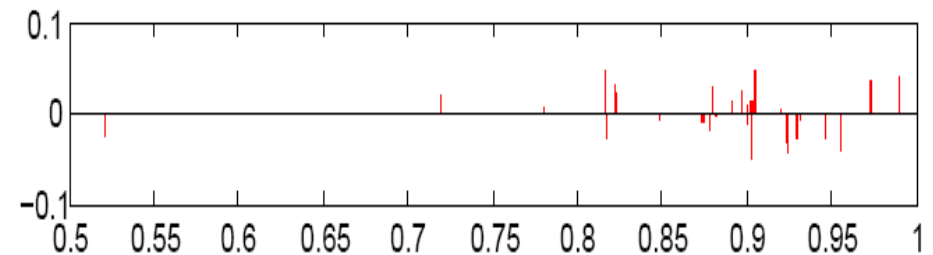
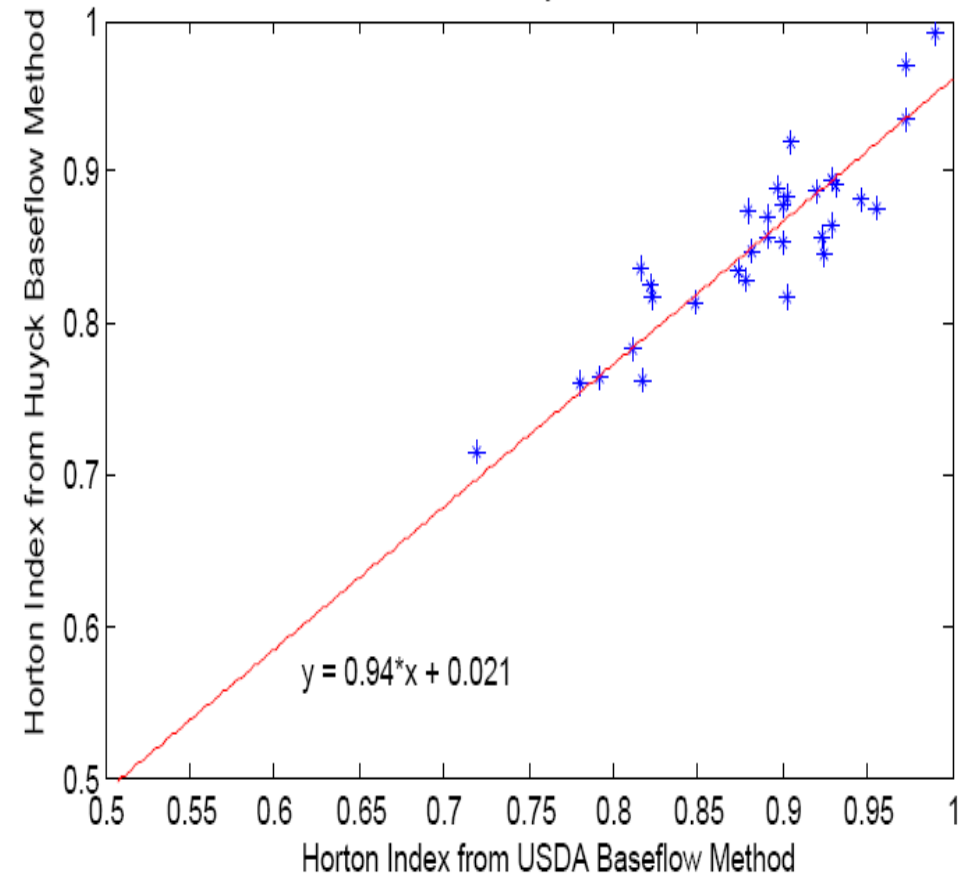
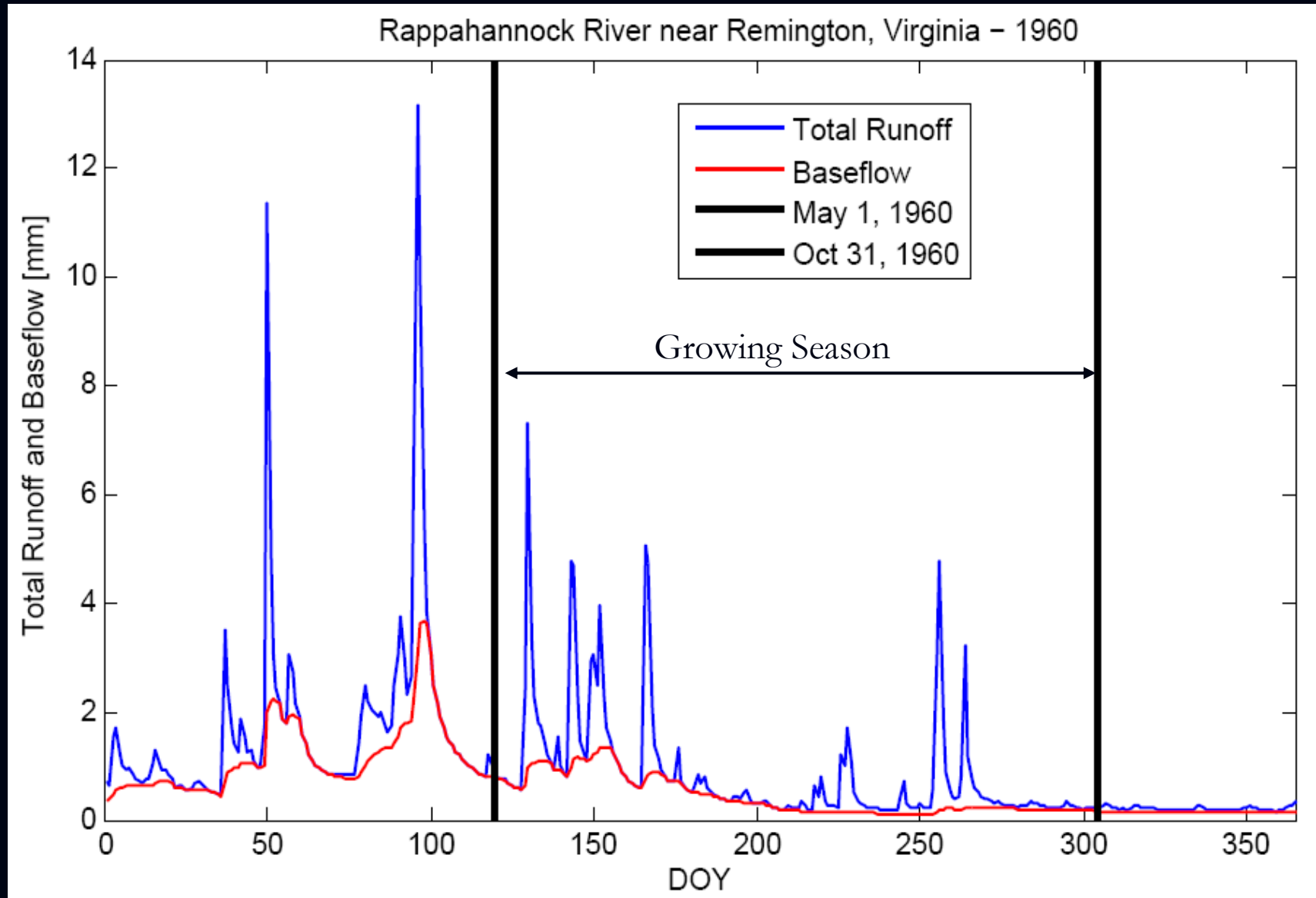
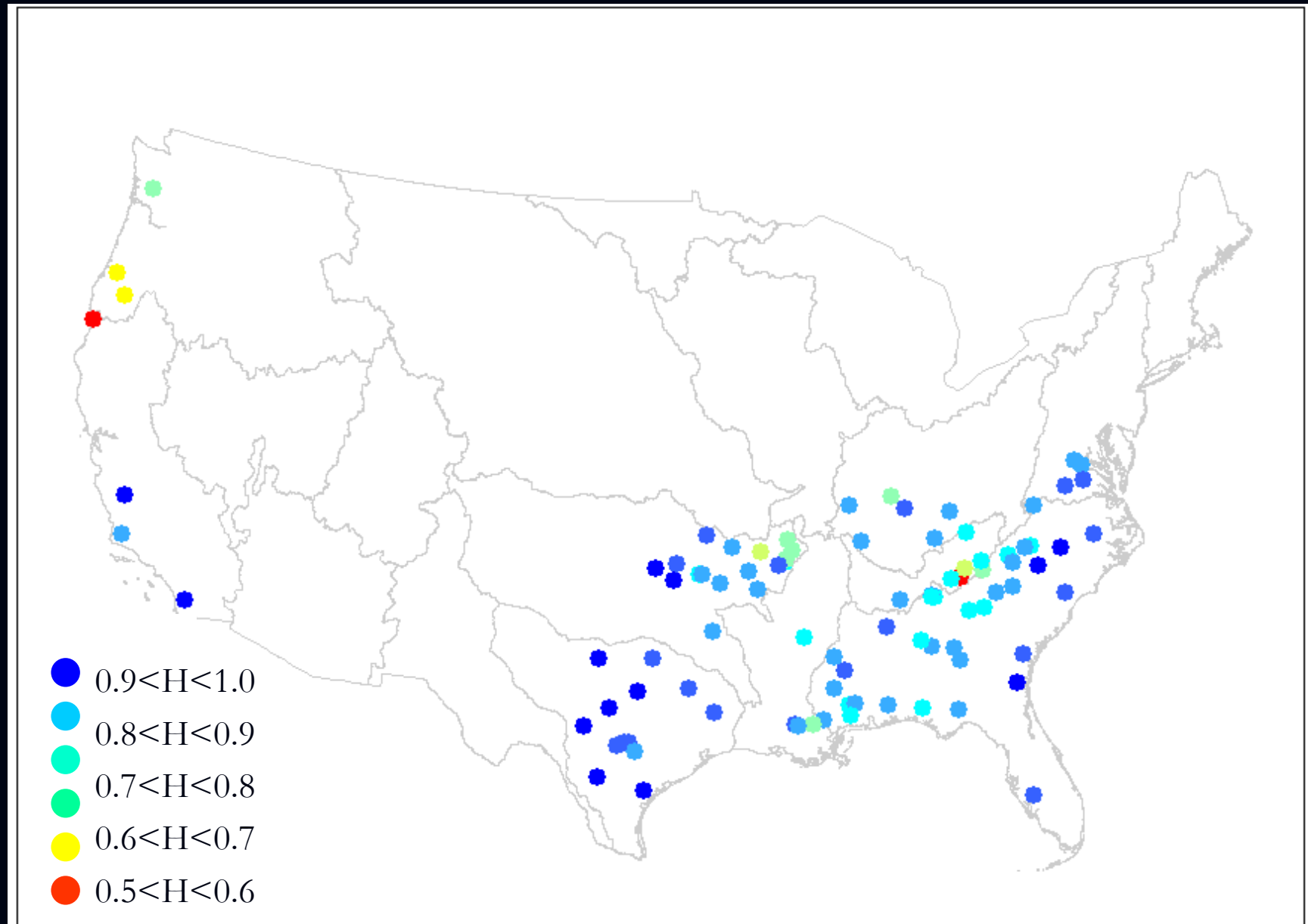


Illustration of *Huyck et al.* Method



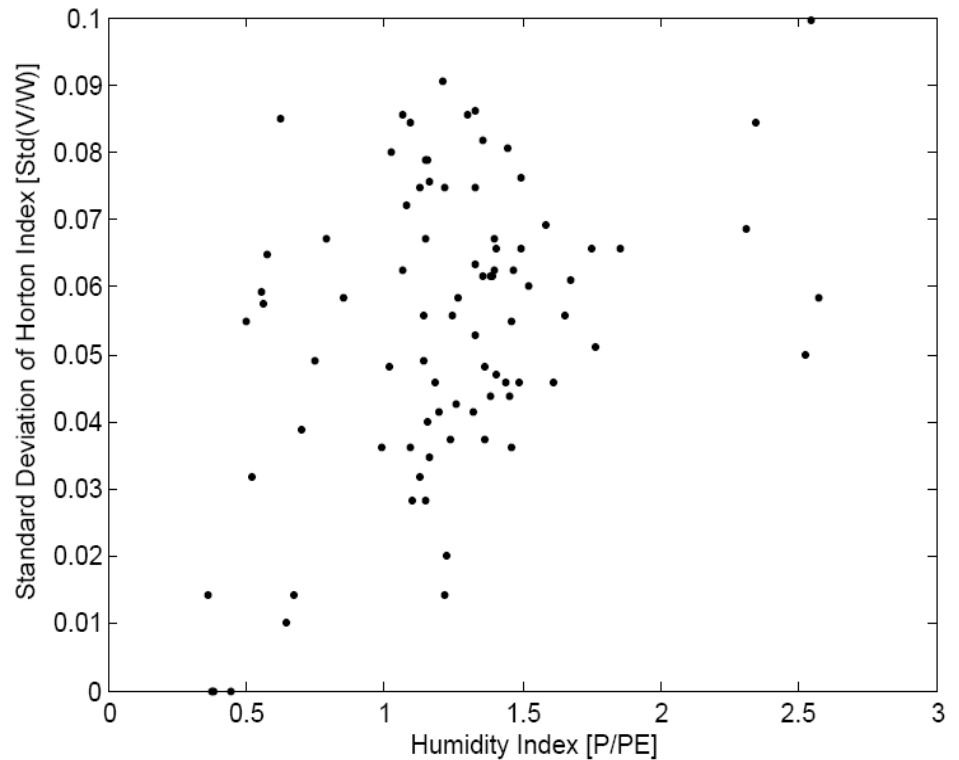
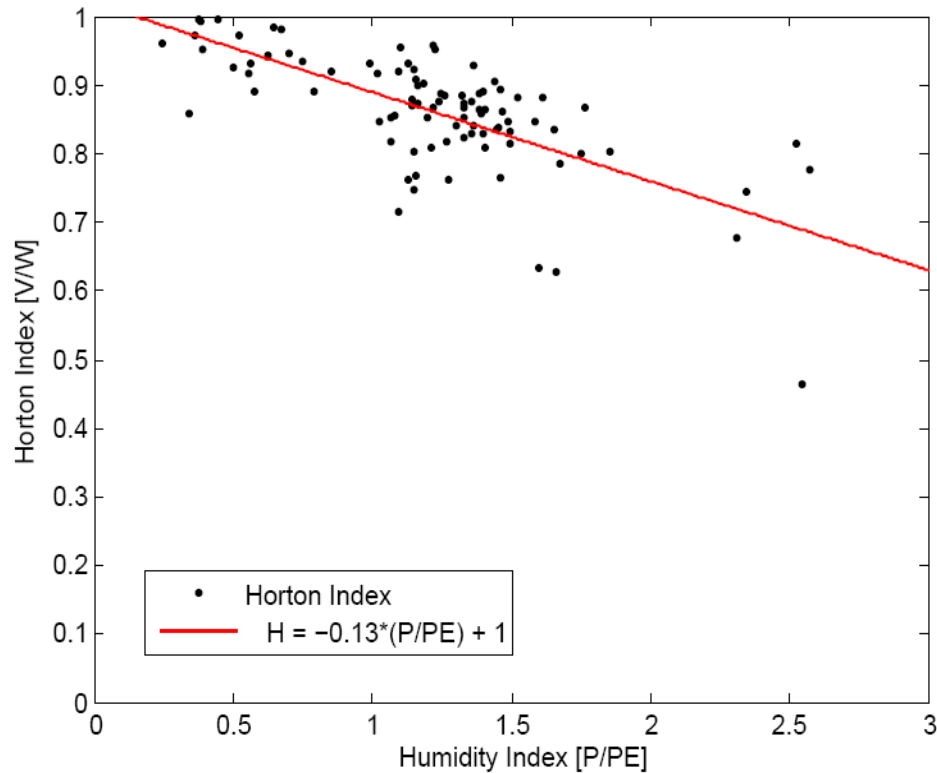
Spatial Variability of Horton Index



Horton Index vs. Humidity Index

Mean Horton Index

Std. Horton Index



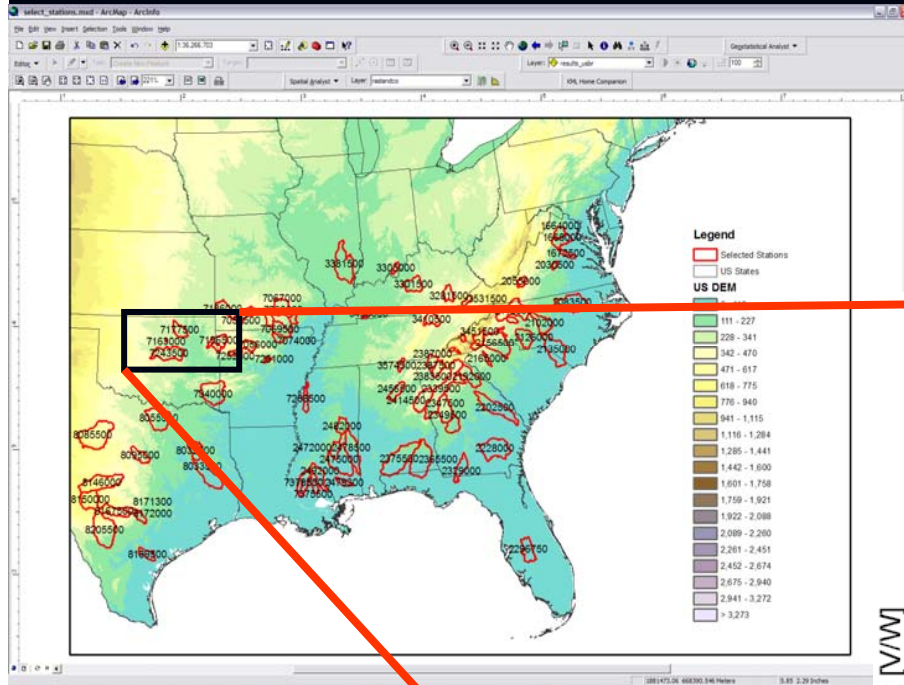
53% with $\text{Std}(H) < 0.06$

74% with $\text{Std}(H) < 0.07$

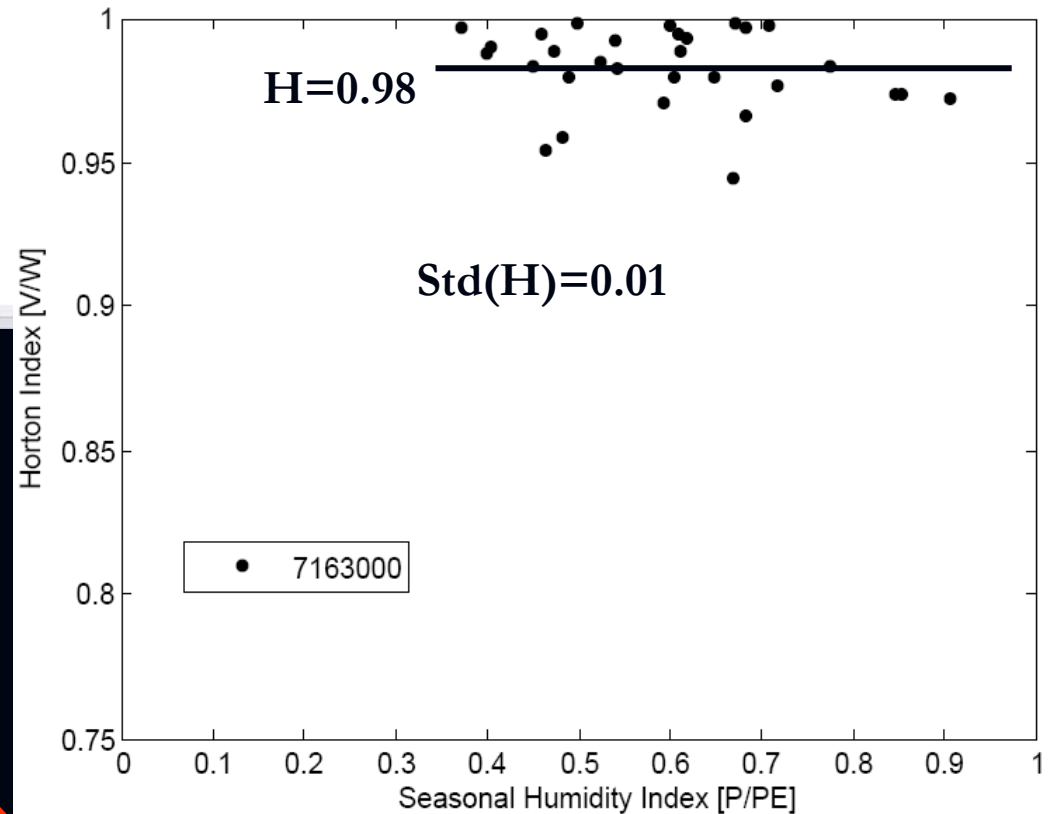
83% with $\text{Std}(H) < 0.08$

93% with $\text{Std}(H) < 0.10$

Interannual Variability of Horton Index



Horton Index Interannual Variability
Savannahs



Ecological controls to interannual variability in semi-arid regions

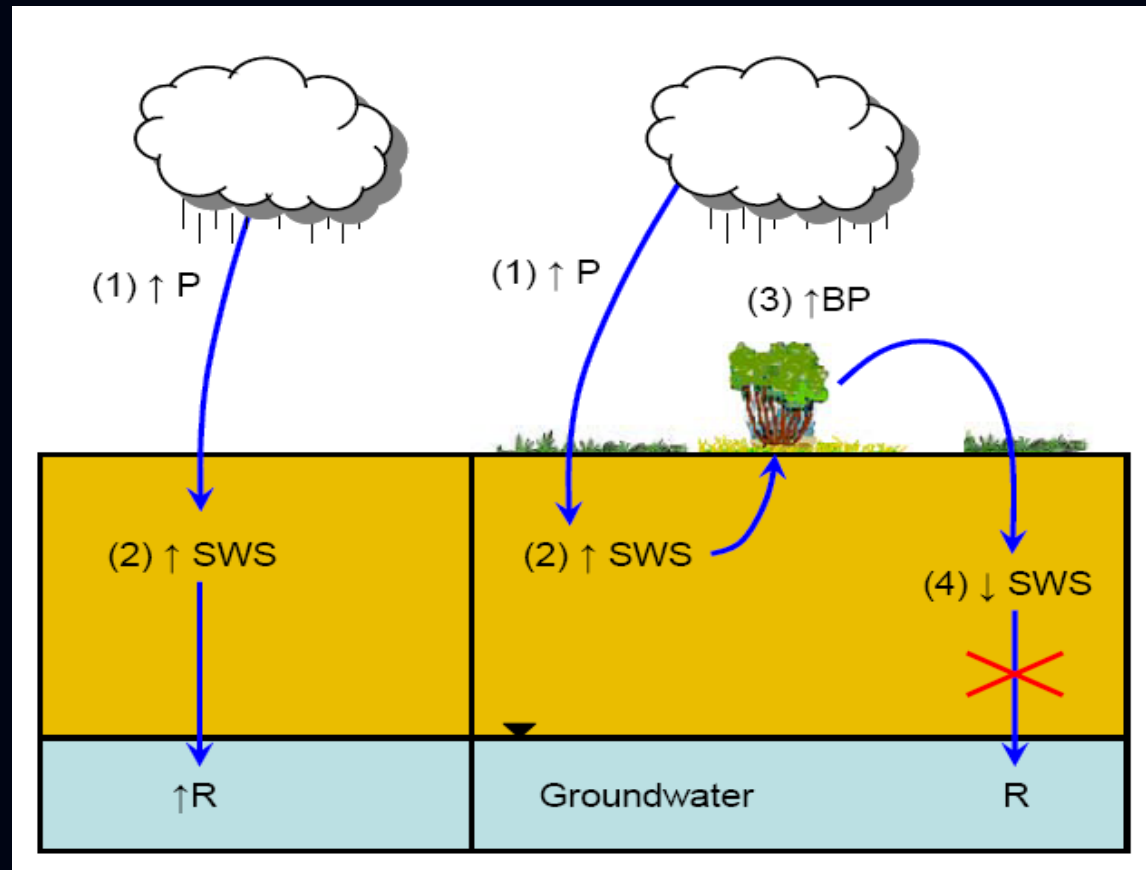
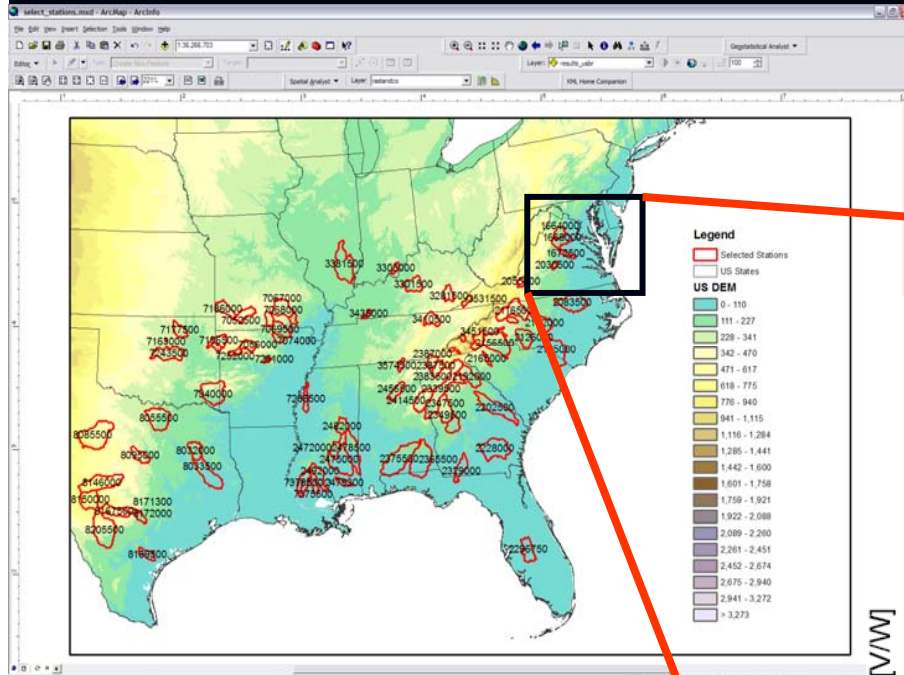
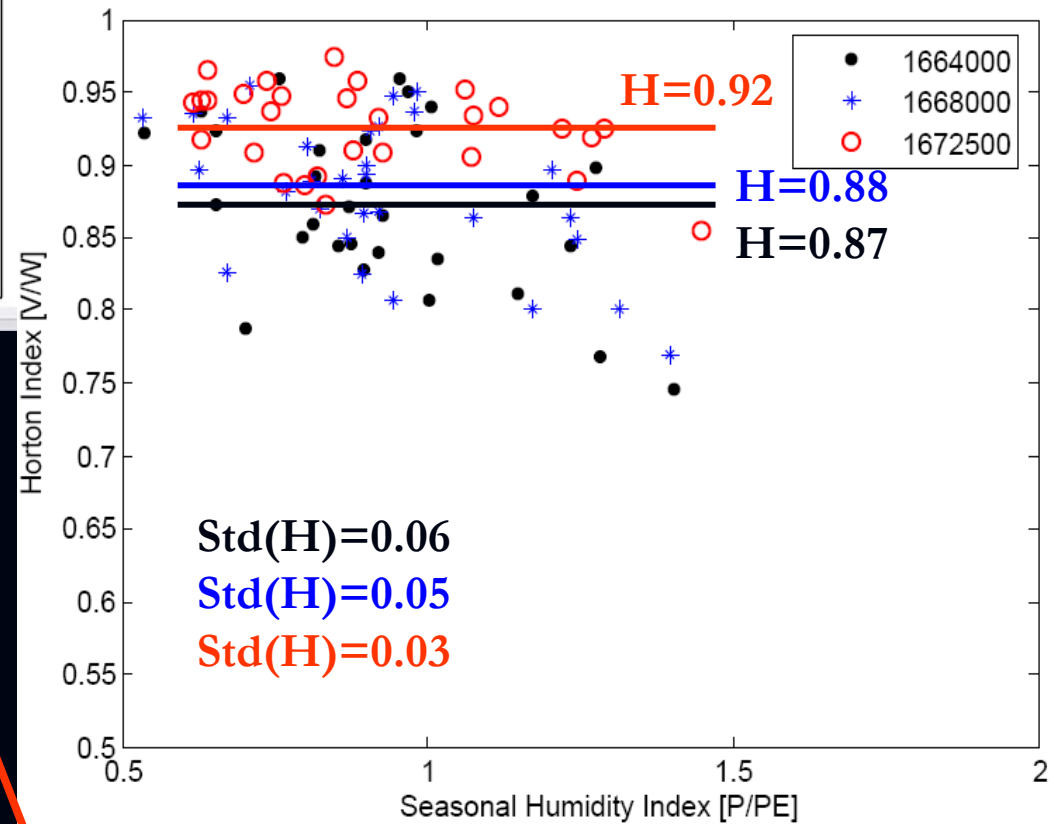


Figure 24: Schematic of non-vegetated and vegetated system responses to elevated precipitation. In non-vegetated systems (*Left*), elevated precipitation (P) results in increased soil-water storage (SWS) that drains resulting in groundwater recharge (R). In the vegetated systems (*Right*), elevated precipitation results in increased soil-water storage that enhances vegetation biomass production (BP), which feeds back to decrease soil-water storage and precludes recharge (Scanlon et al., 2005).

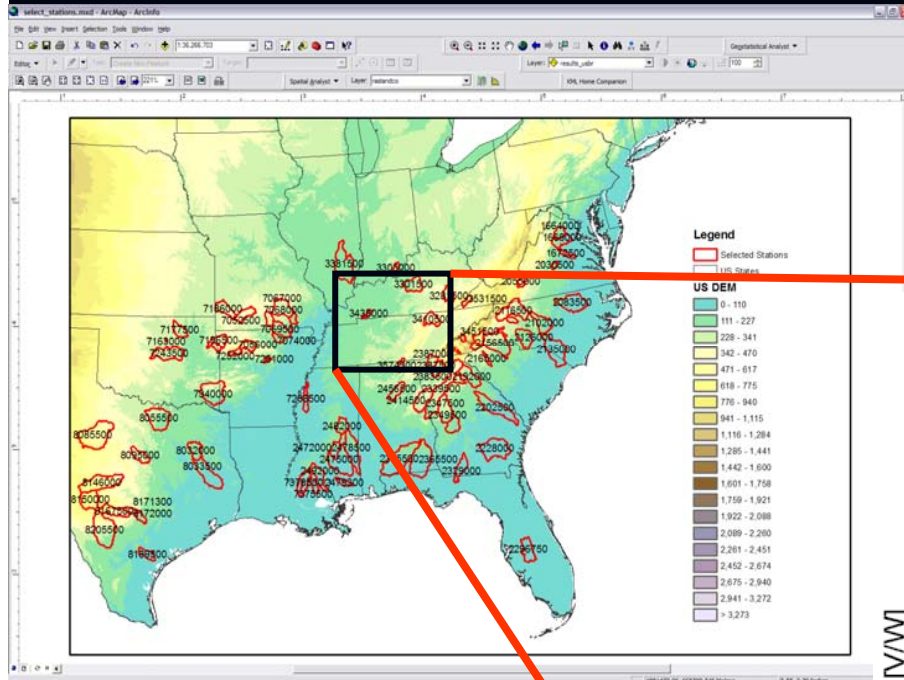
Interannual Variability of Horton Index



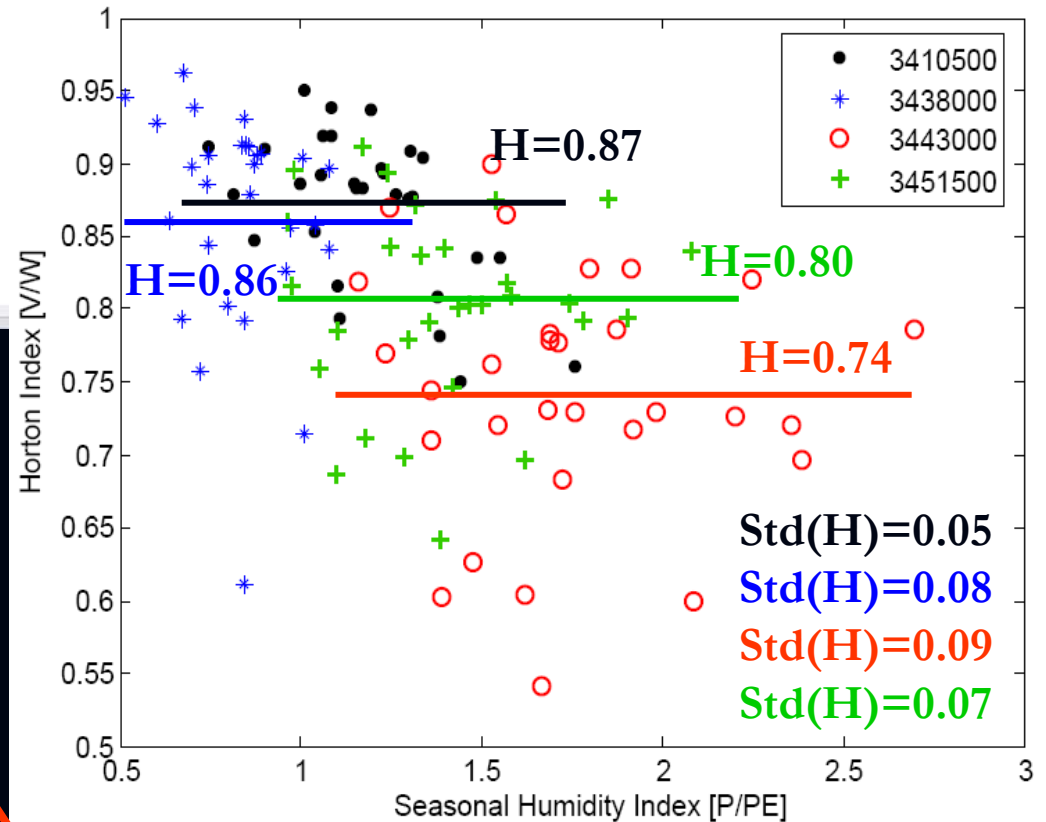
Horton Index Interannual Variability
Closed Shrublands



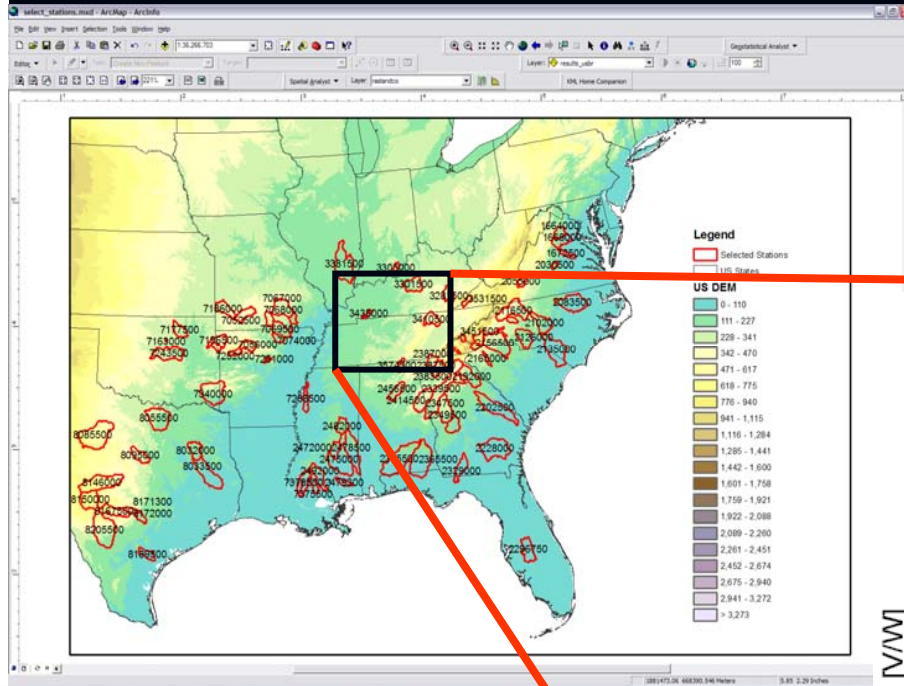
Interannual Variability of Horton Index



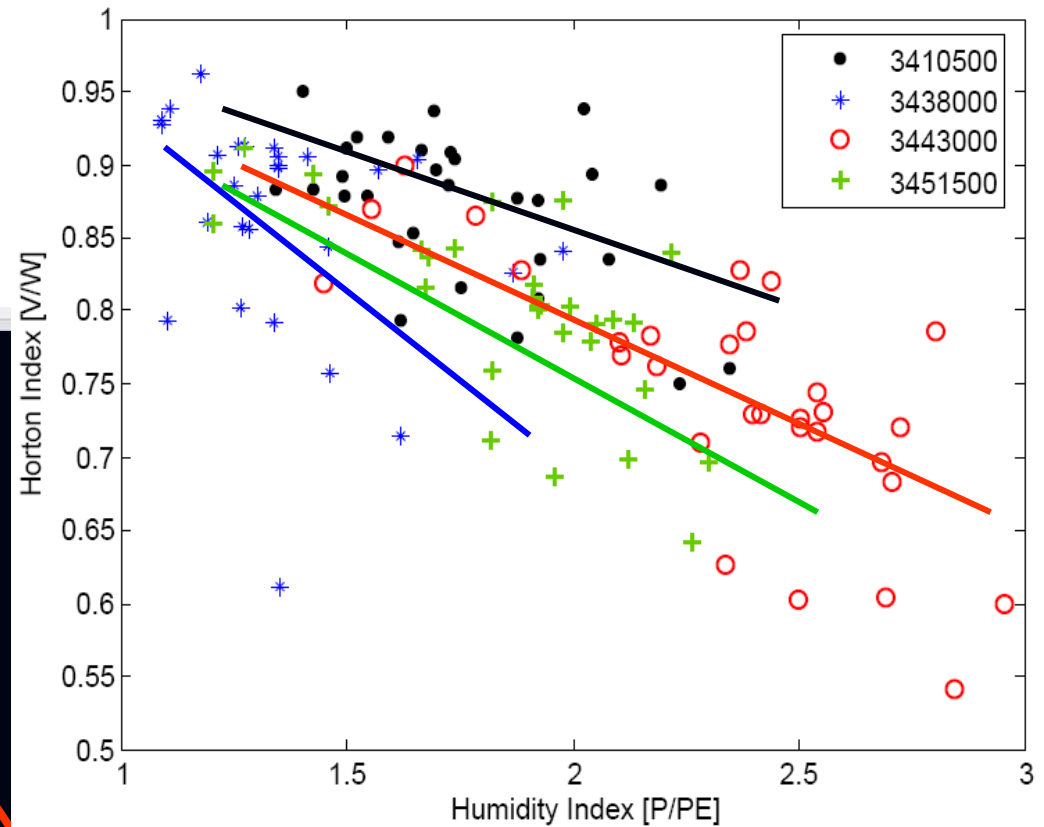
Horton Index Interannual Variability
Closed Shrublands



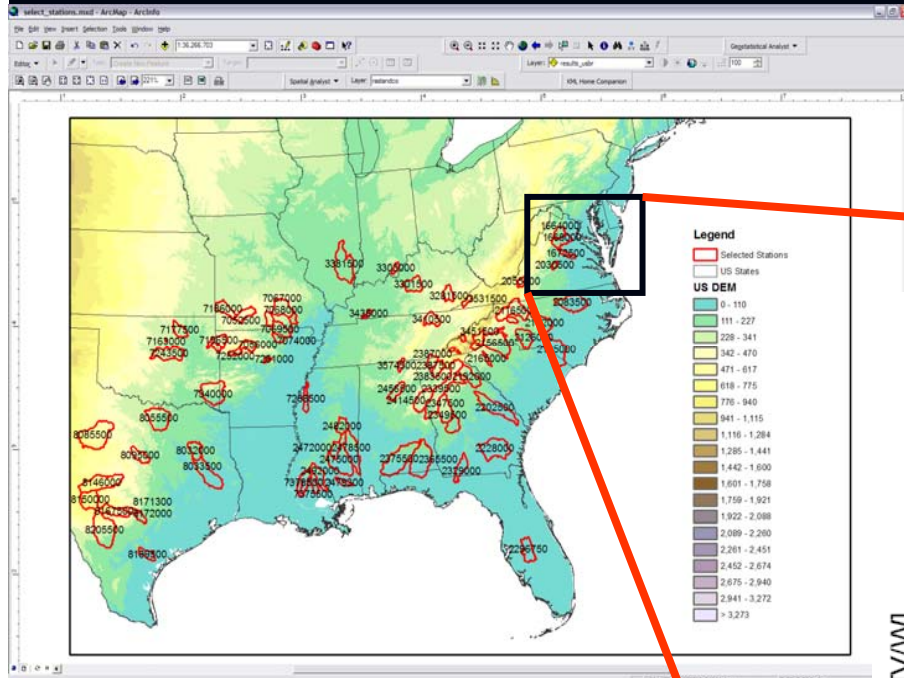
Interannual Variability of Horton Index



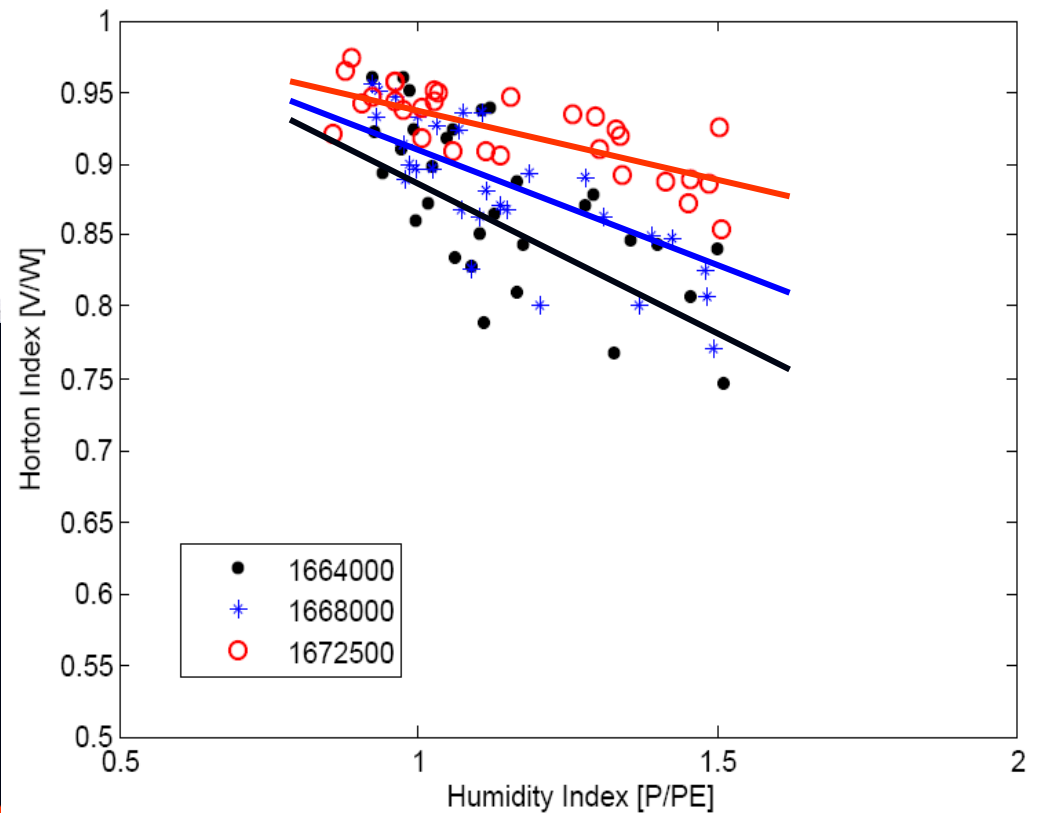
Horton Index Interannual Variability
Closed Shrublands



Interannual Variability of Horton Index



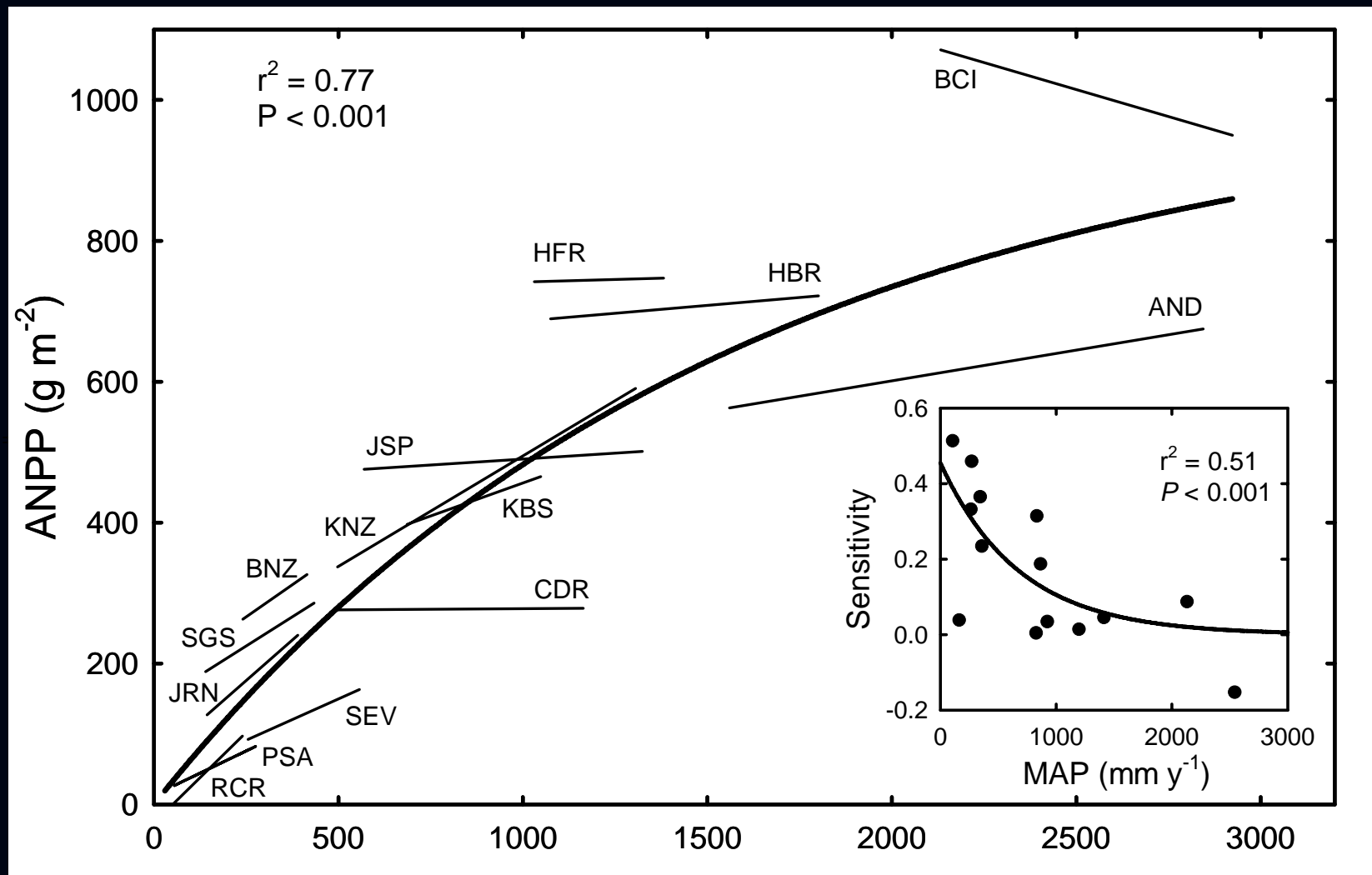
Horton Index Interannual Variability
Closed Shrublands



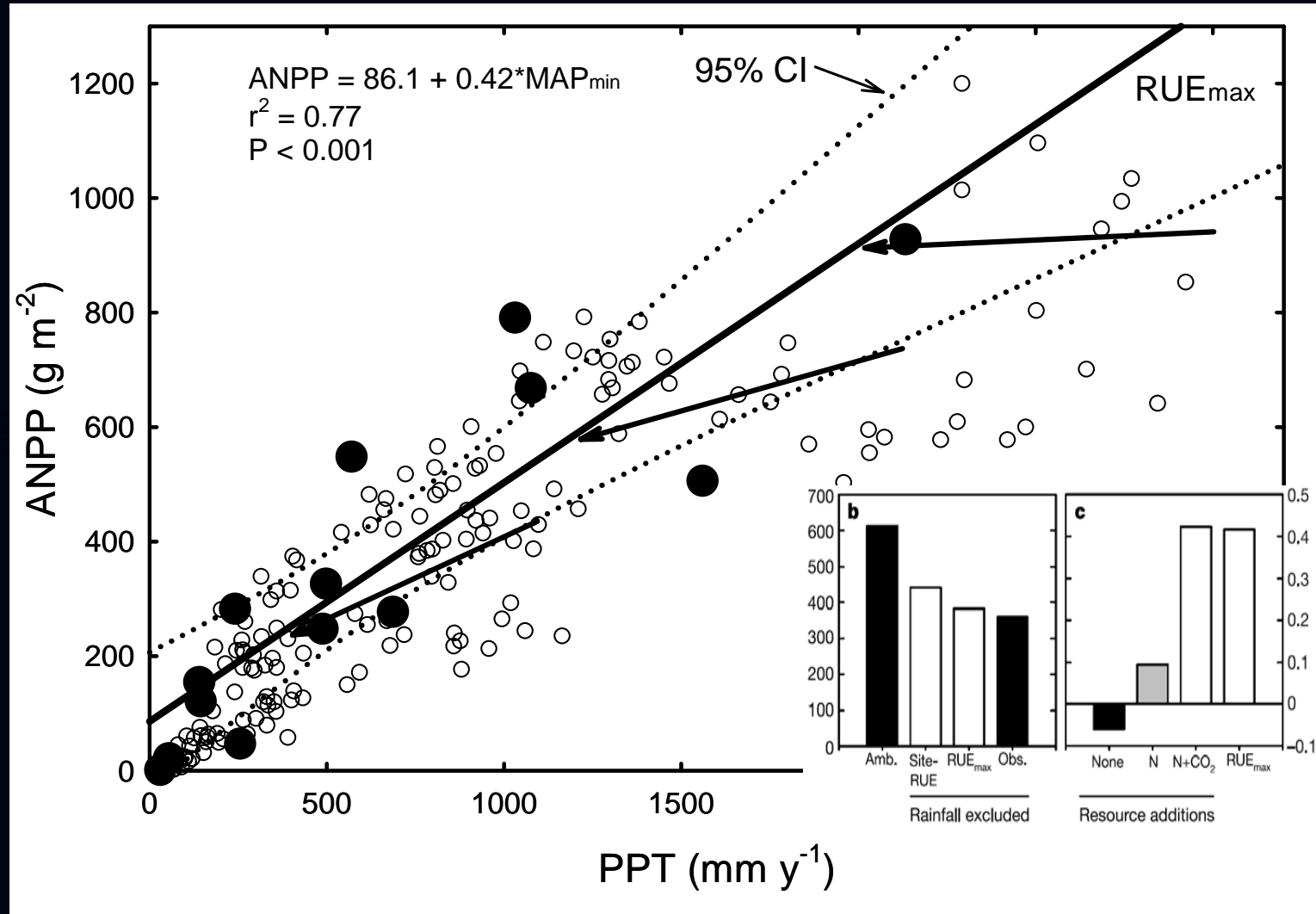
Precipitation and production



Biome rainwater use efficiency



Convergence to a common RUE_{max}



Water Use Efficiency and Actual ET

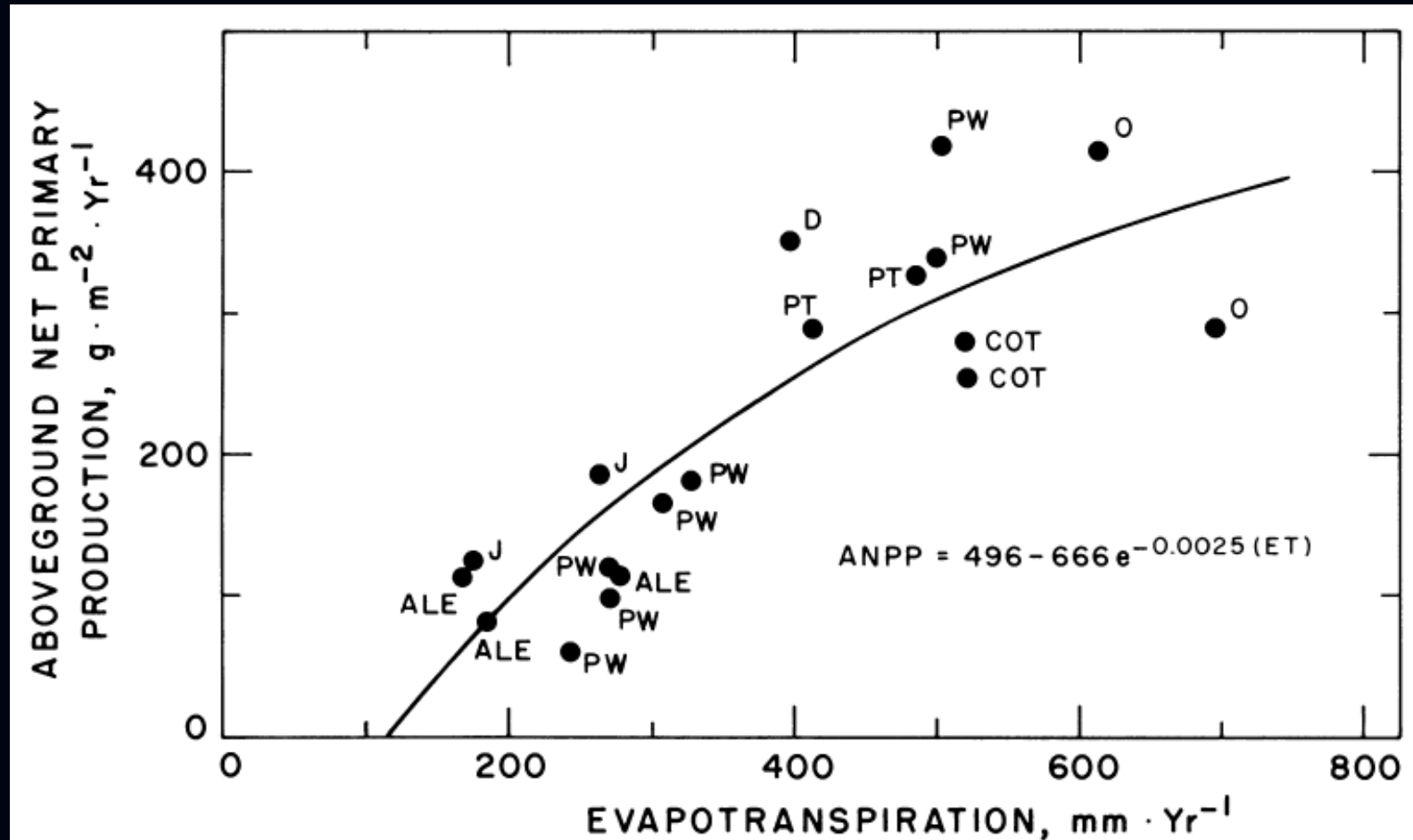
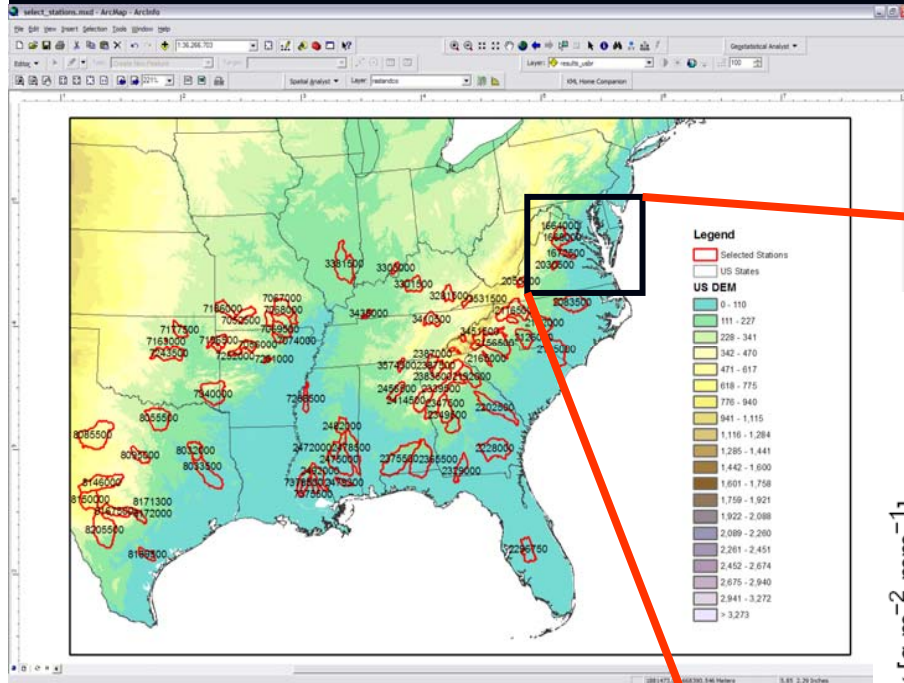
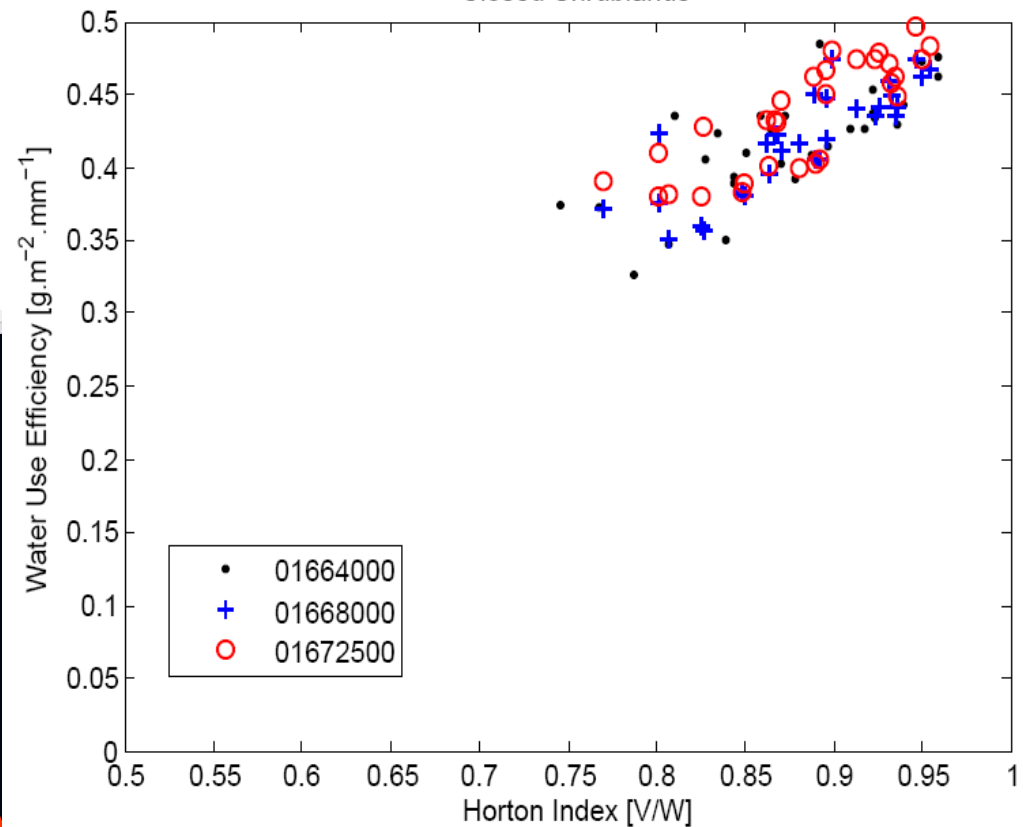


FIG. 2. Aboveground net primary production (ANPP) and actual evapotranspiration (AET) from several United States grasslands sites fully described by Sims et al. (1978). Location and dominant vegetation are listed in Table 1. Site abbreviations are ALE, Ale, Washington, USA; COT, Cottonwood, South Dakota, USA; D, Dickinson, South Dakota, USA; J, Jornada, New Mexico, USA; O, Osage, Oklahoma, USA; PW, Pawnee, Colorado, USA; and PT, Pantex, Texas, USA.

Catchment-scale Water Use Efficiency



Catchment-scale Water Use Efficiency
Closed Shrublands



$$\text{ANPP} = f(V)$$

$$\text{WUE} = \frac{\text{ANPP}}{P - S}$$

The annual water balance

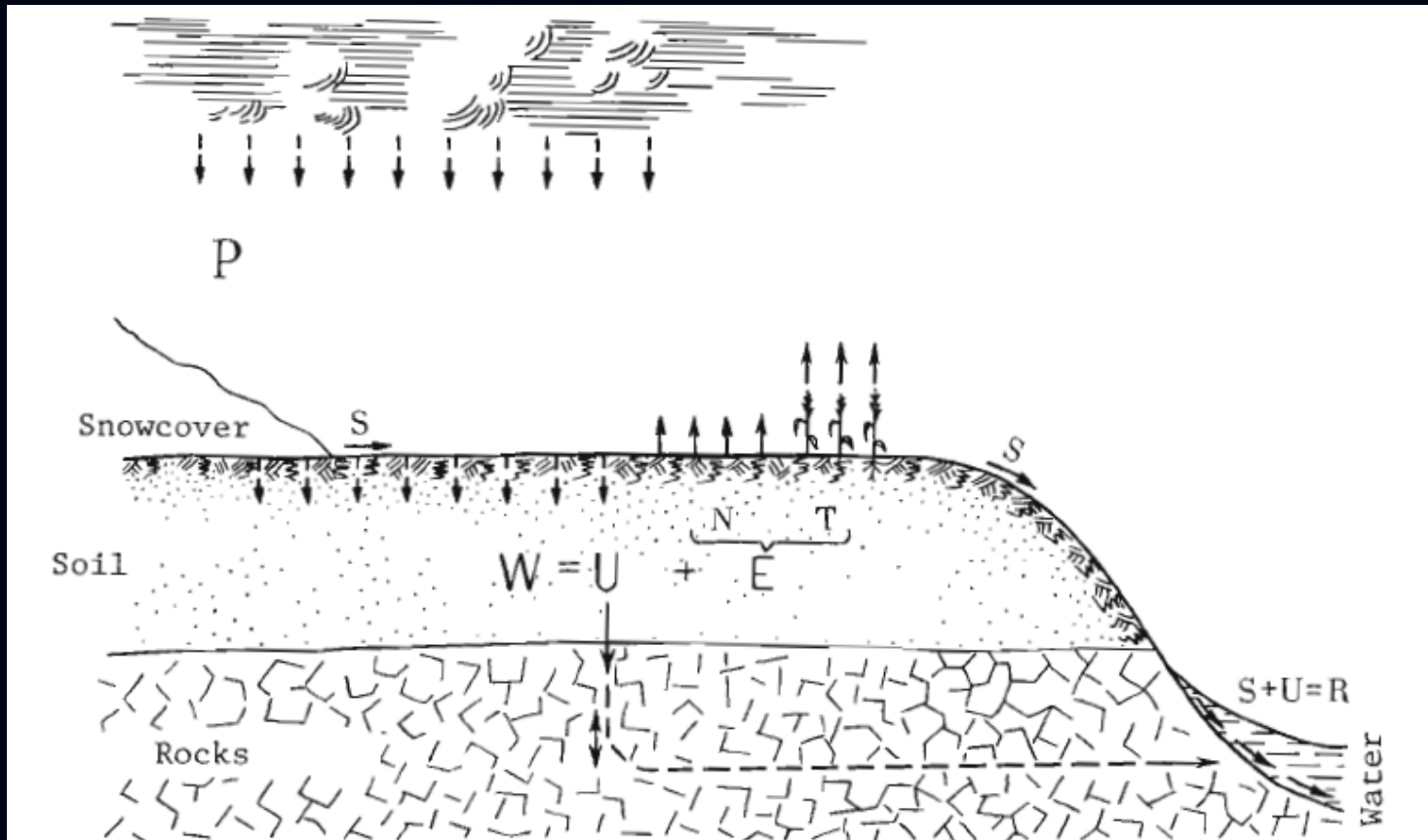
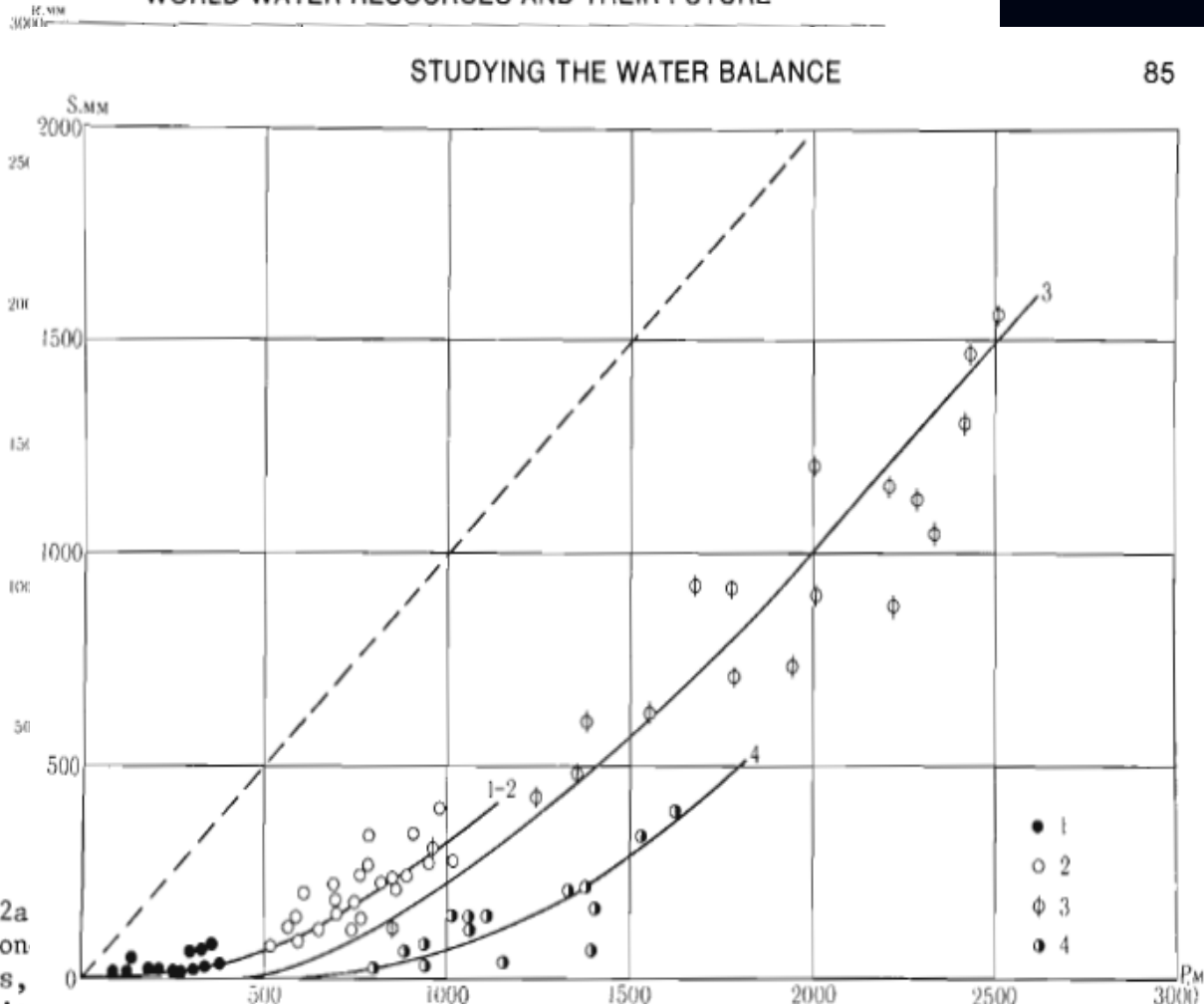


Fig. 3. Diagram of the water balance of land area. P--Precipitation; R--total runoff; U--groundwater runoff; S--surface runoff; W--total wetting of the area (annual infiltration) including surface retention; N--unproductive evaporation (evaporation proper); T--transpiration of plants; E--evapotranspiration.

The L'vovich Hypothesis

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WORLD WATER RESOURCES AND THEIR FUTURE



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Fig. 12a
function
meadows,
mountain

Fig. 12b. Plot of interpolation formula for surface runoff (S) as a function of rainfall (P) for South America. Notations same as in Fig. 12a.

$$P = R + V$$

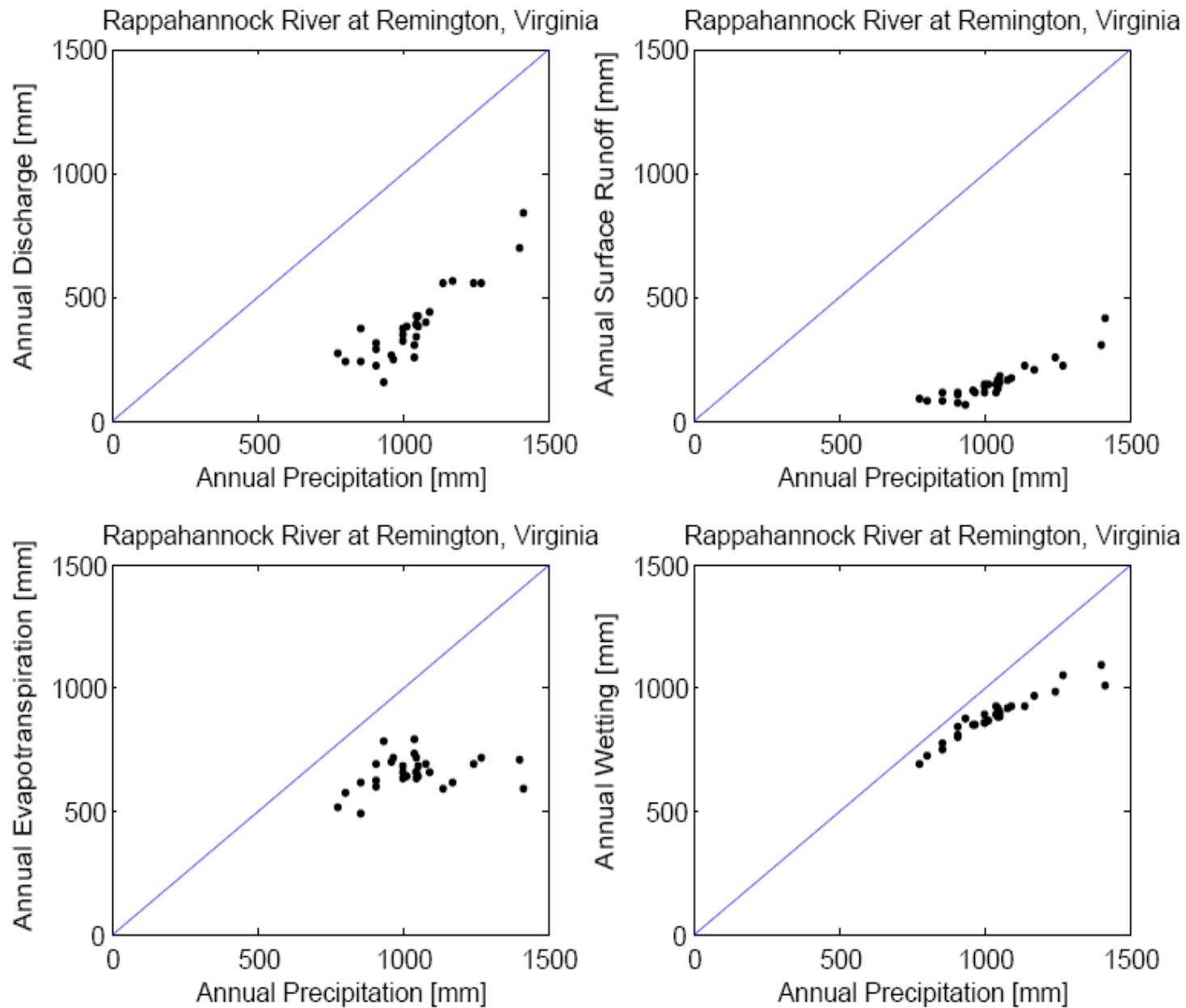
$$P = S + W$$

$$R = S + U$$

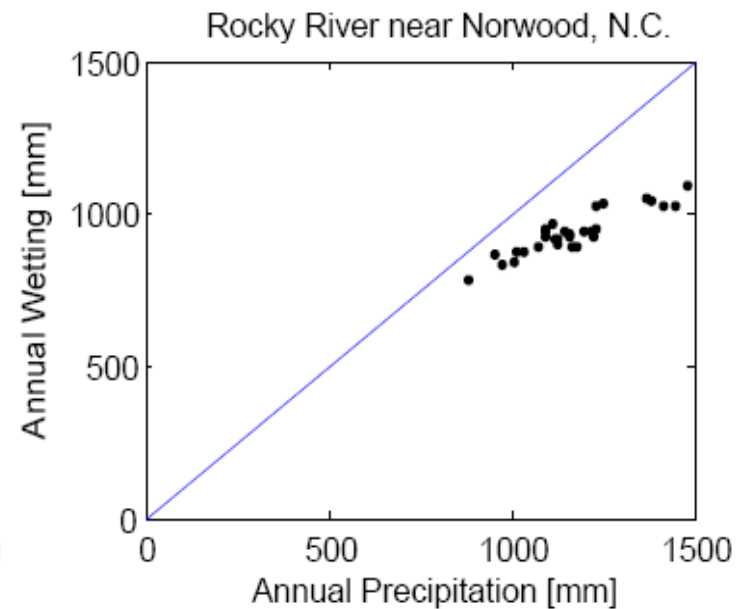
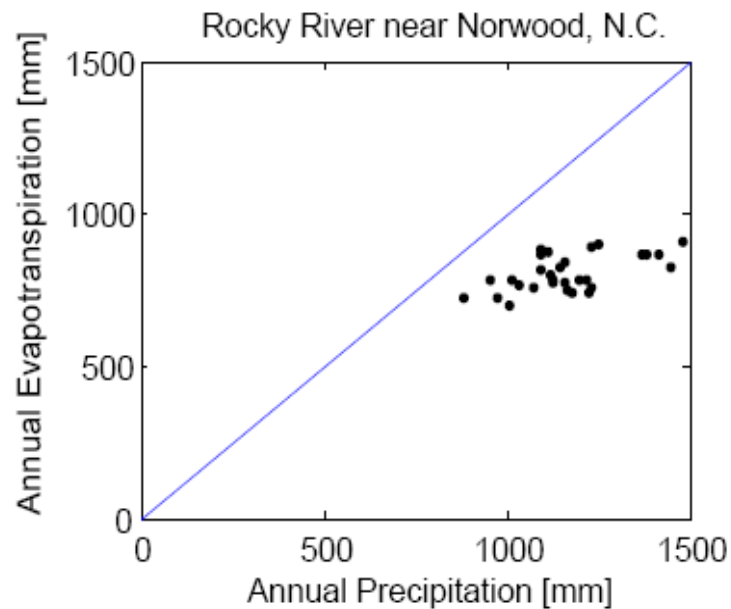
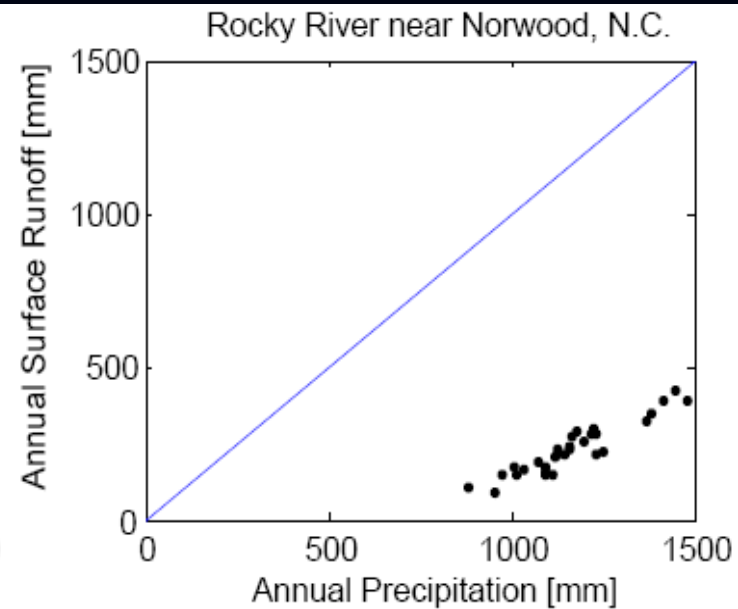
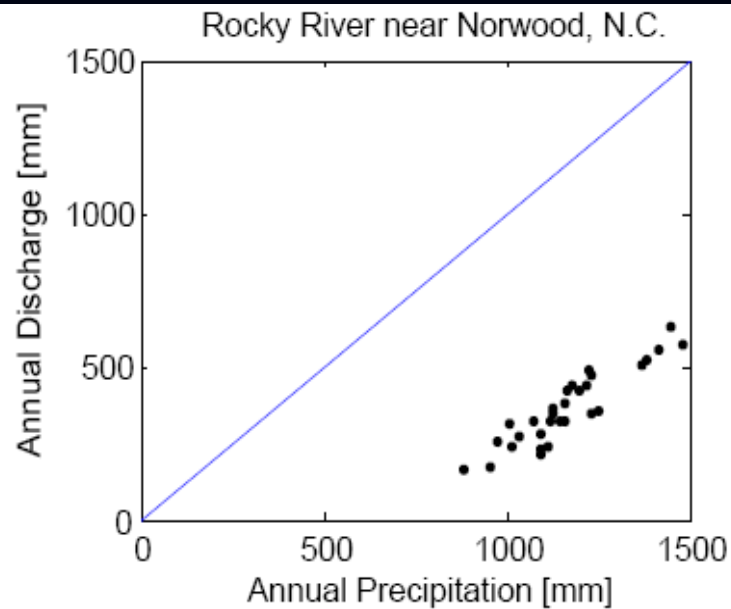
$$W = U + V$$

L'vovich, 1979 (AGU)

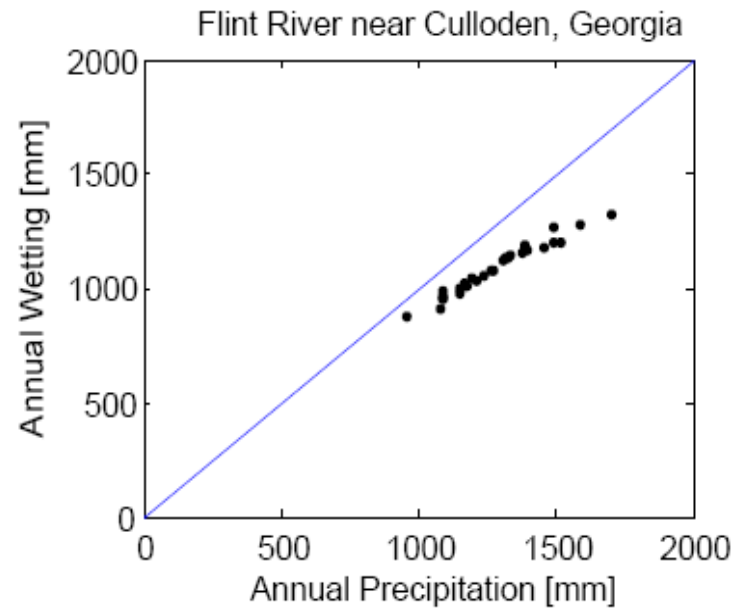
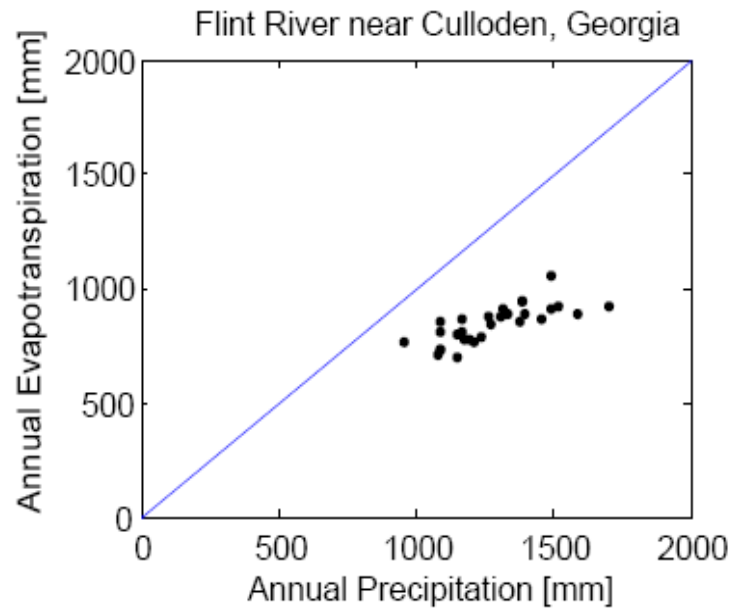
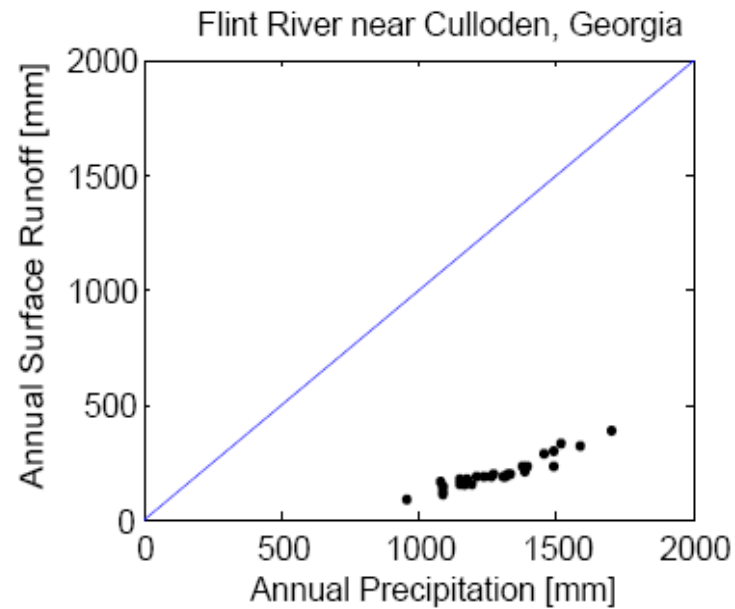
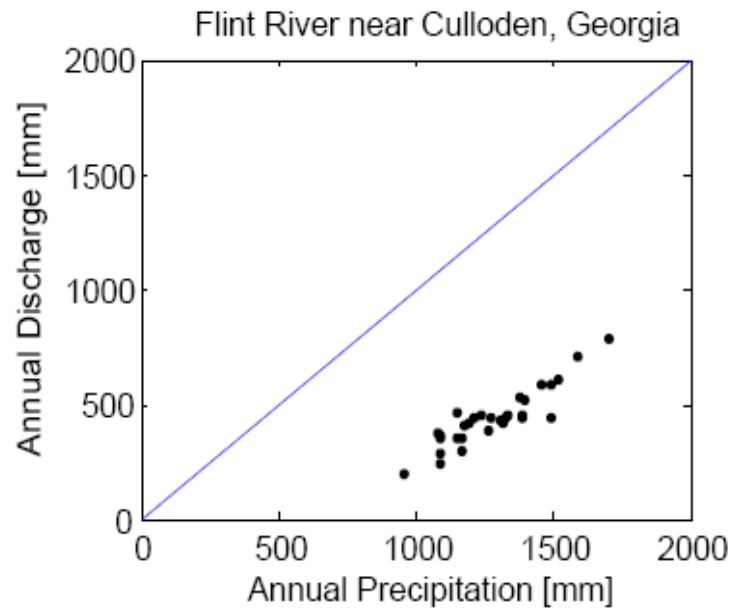
Rappahannock River, Virginia



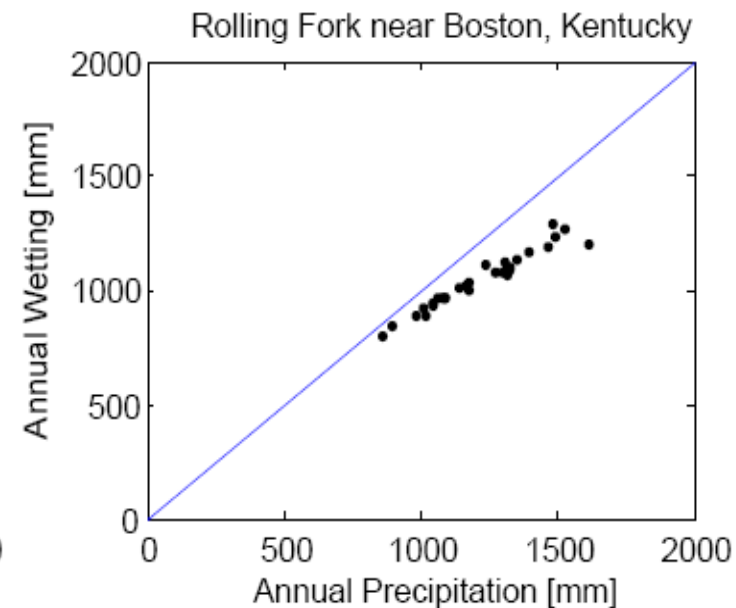
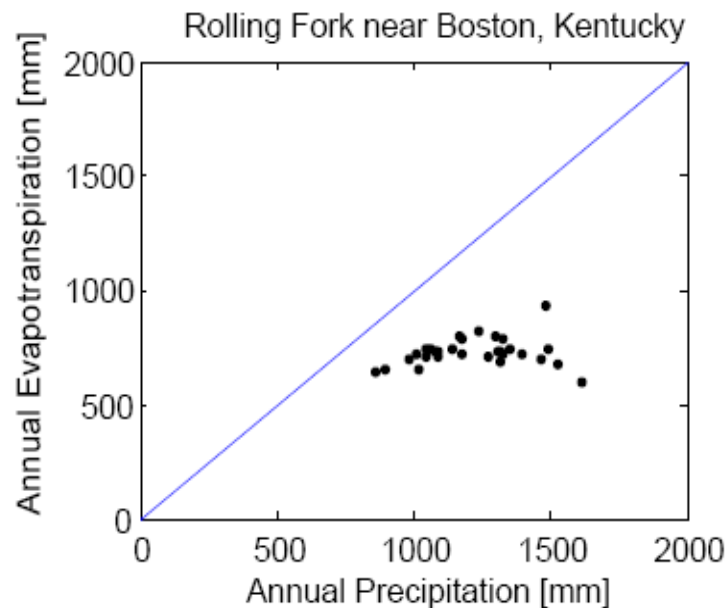
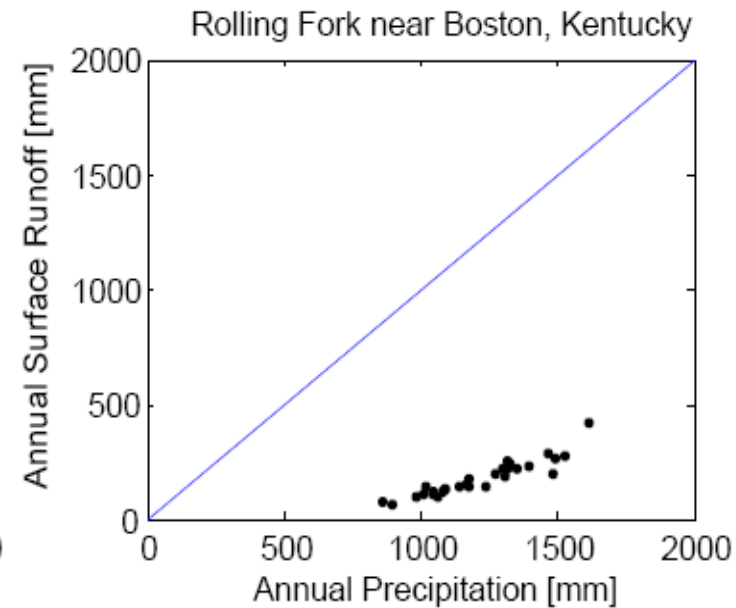
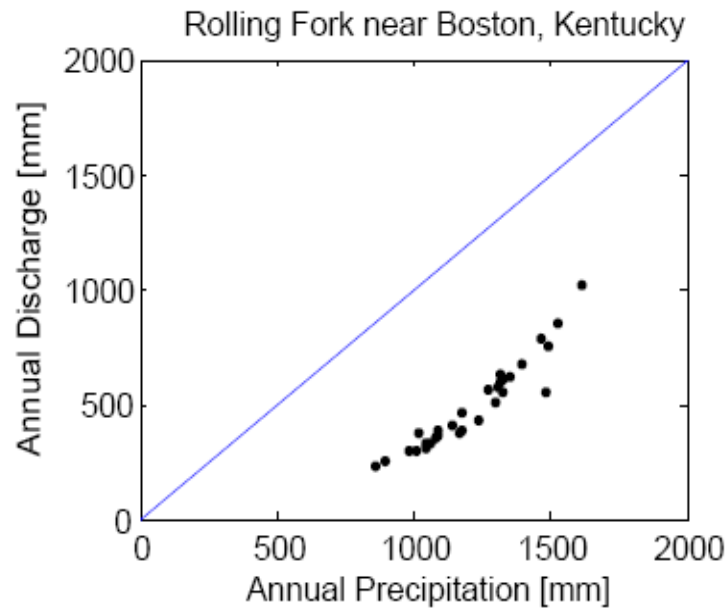
Rocky River, North Carolina



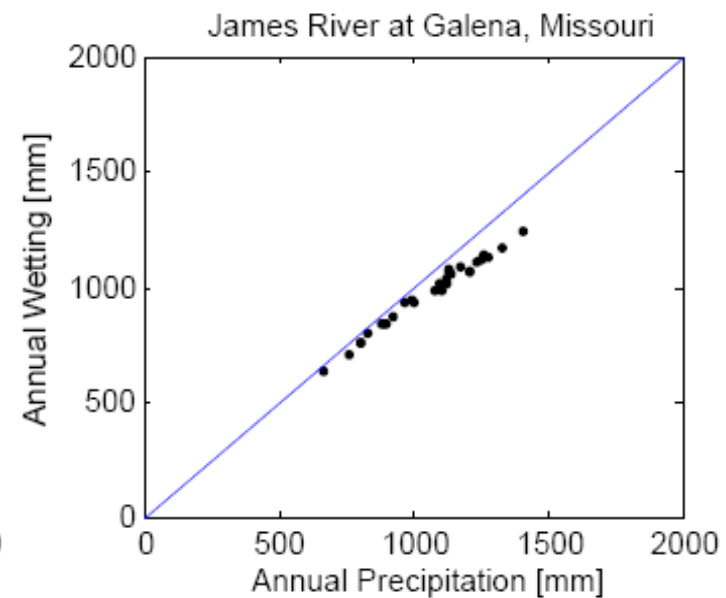
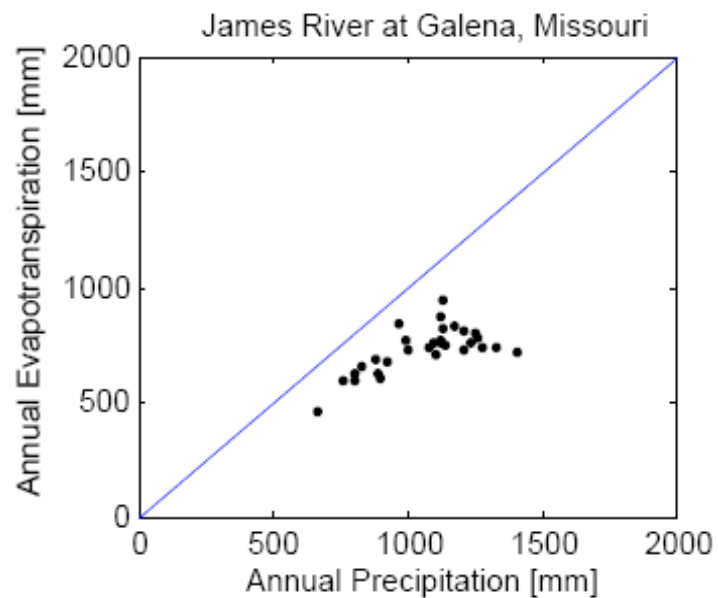
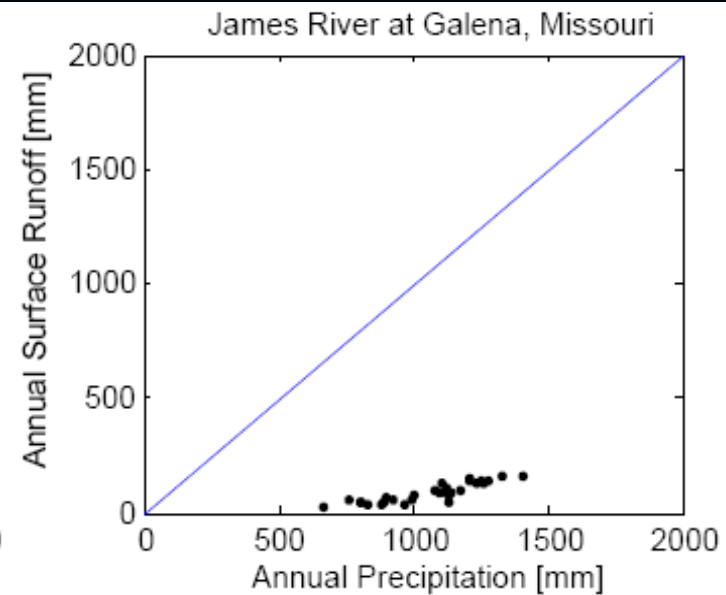
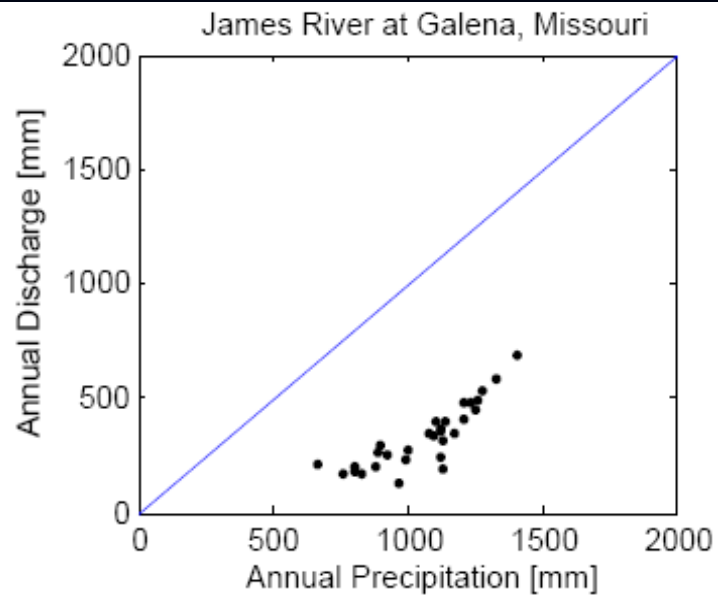
Flint River, Georgia



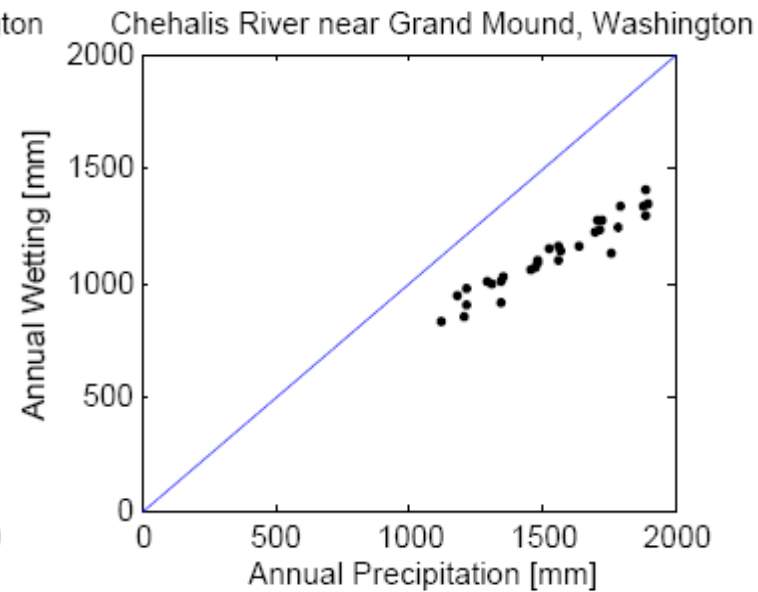
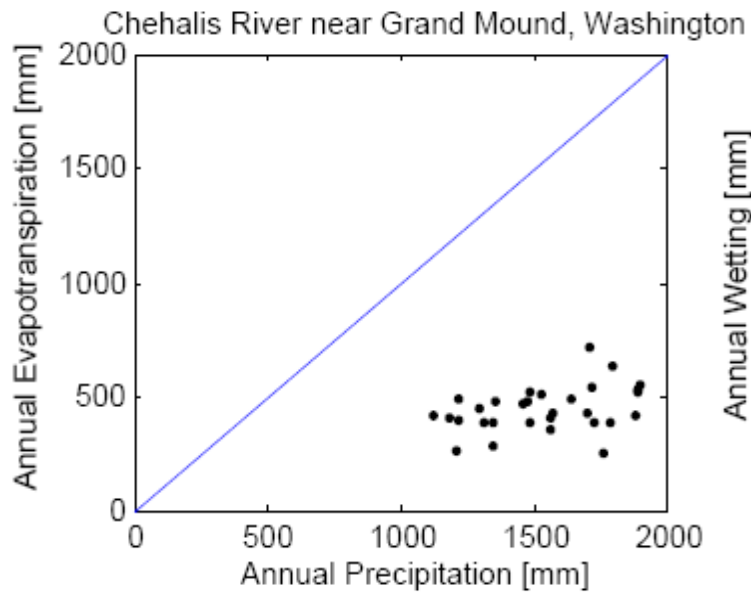
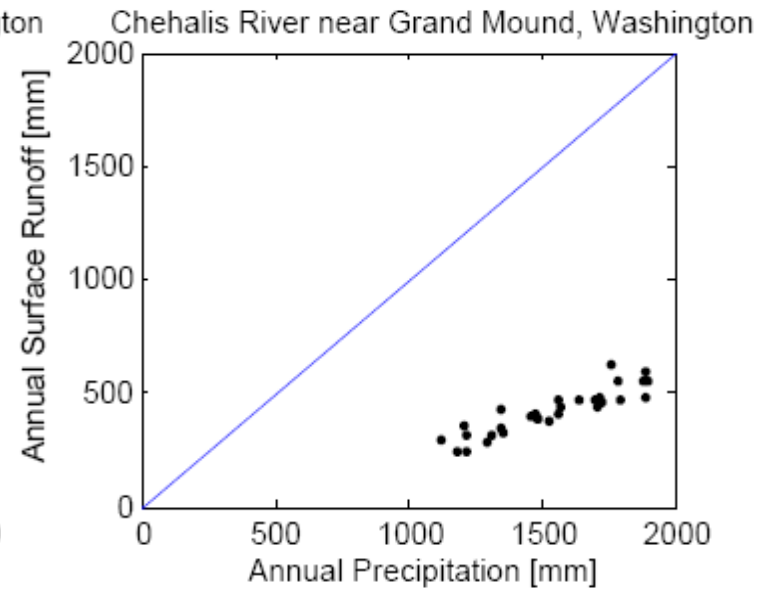
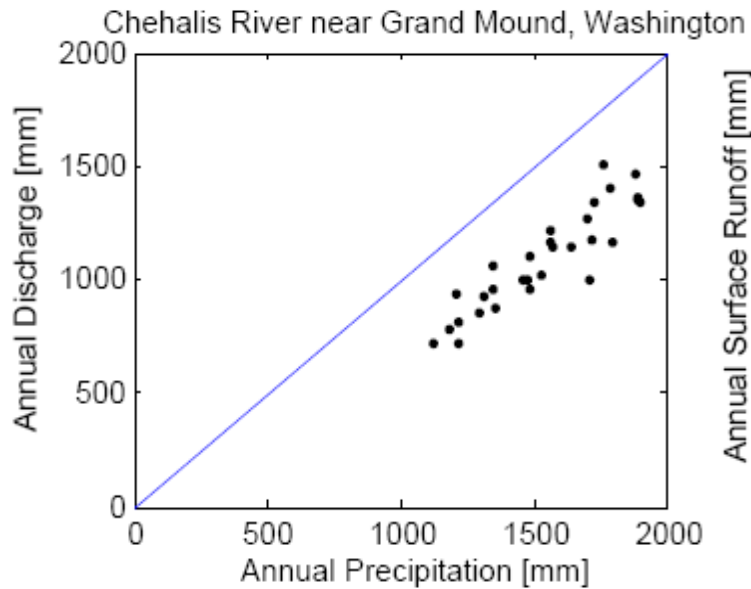
Rolling Fork River, Kentucky



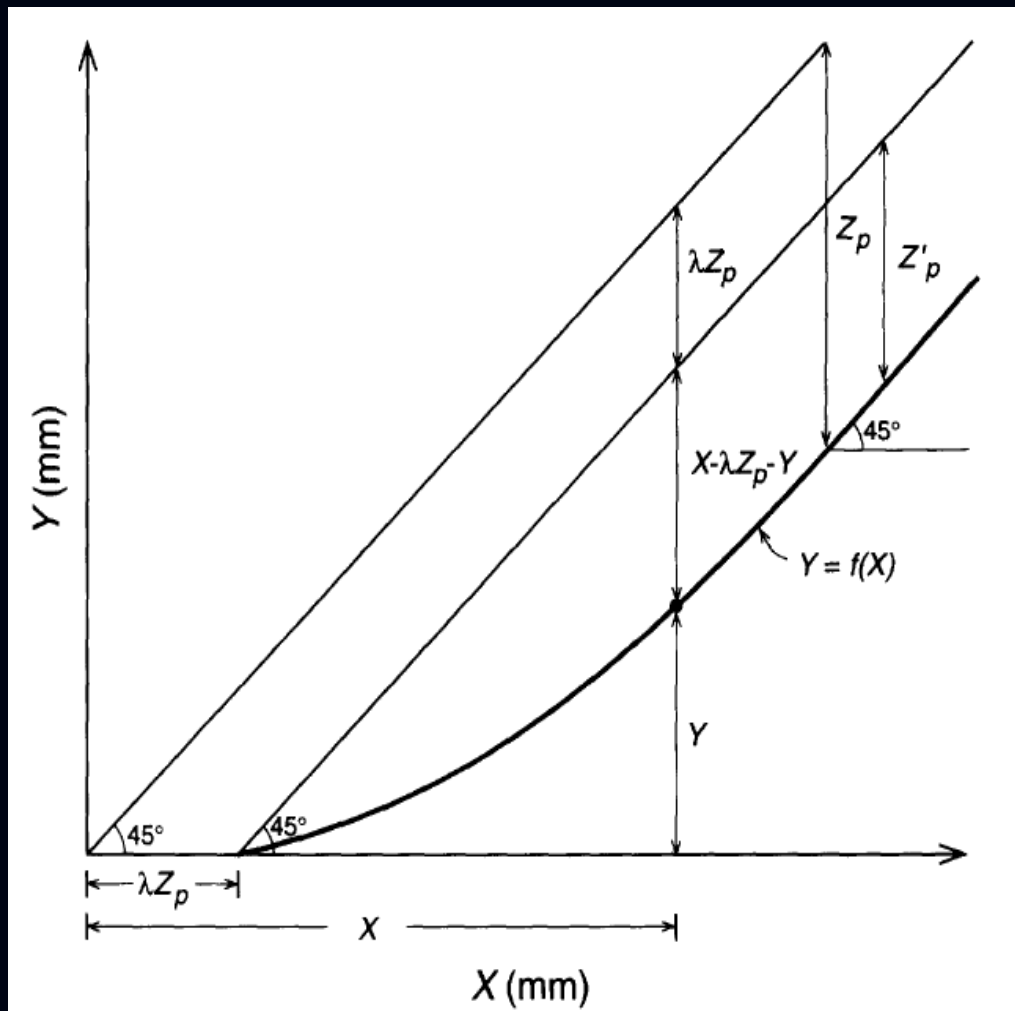
James River, Missouri



Chehalis River, Washington



Proportionality Relations



$$X = Y + Z$$

$$Z = X - Y$$

$$Z \rightarrow Z_p \Leftrightarrow X \rightarrow \infty; Y \rightarrow \infty$$

$$P = S + U + V = S + W$$

$$W = P - S$$

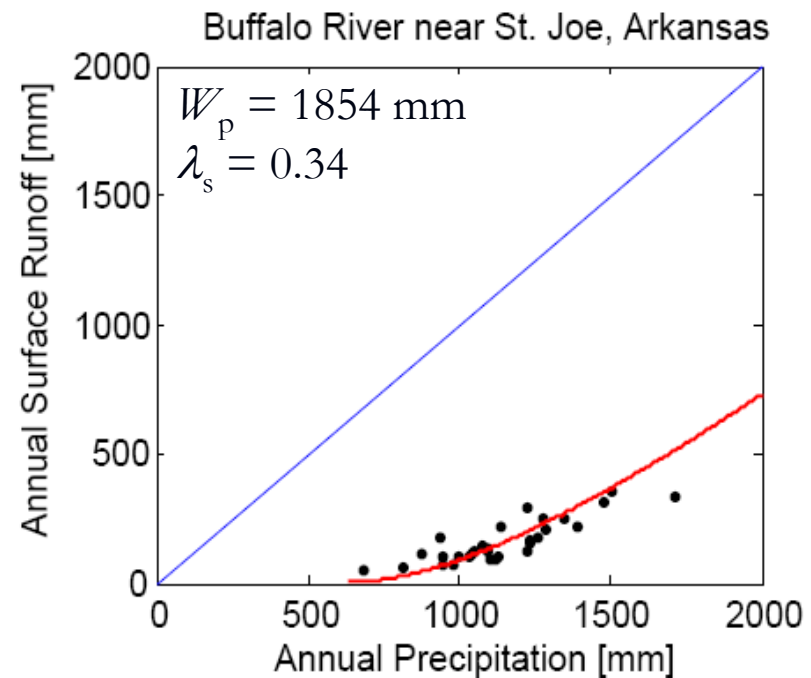
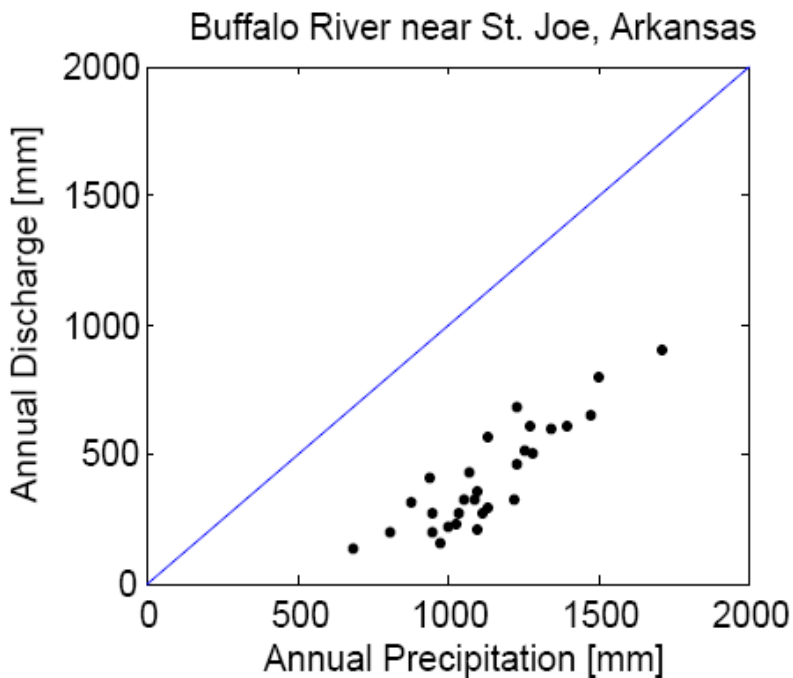
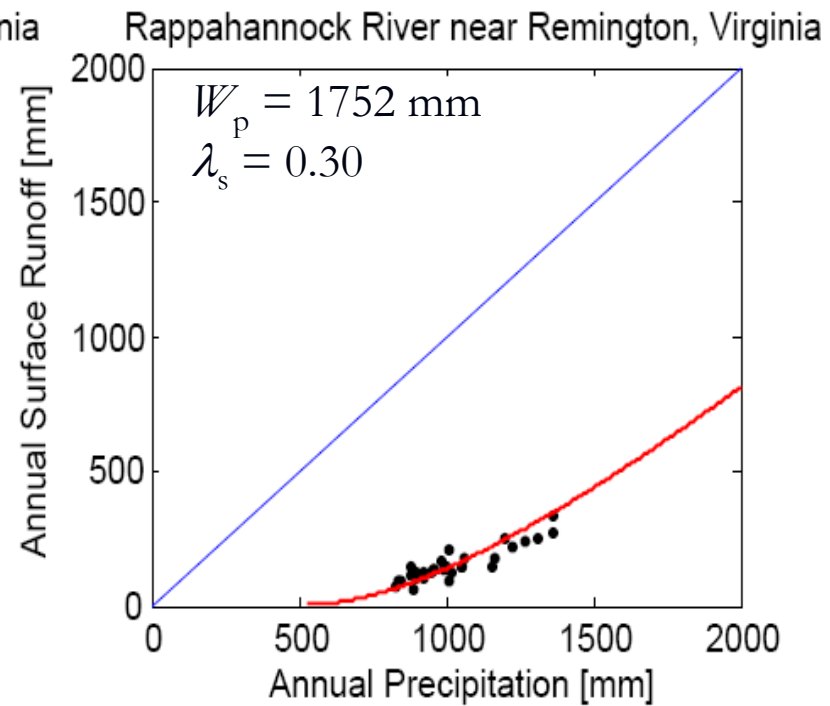
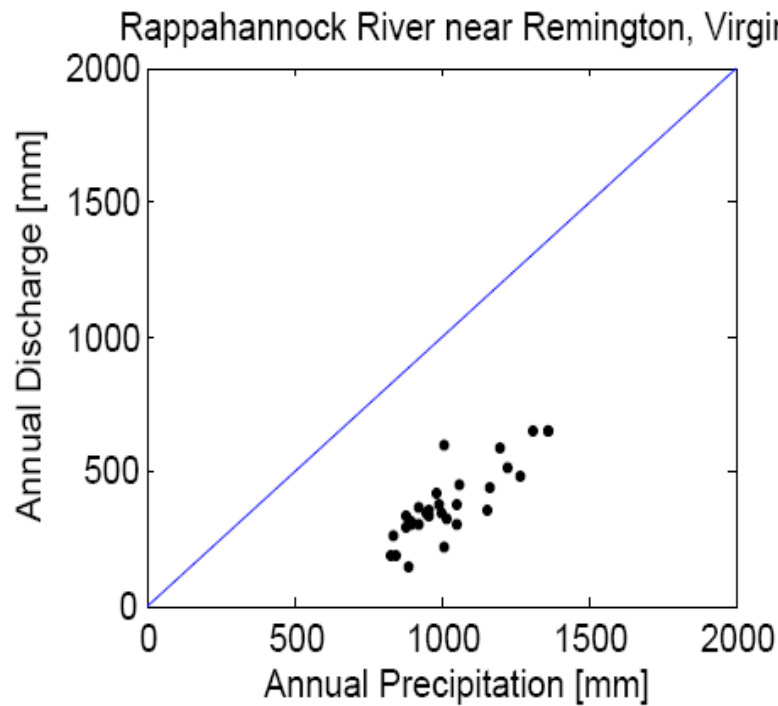
$$W \rightarrow W_p \Leftrightarrow P \rightarrow \infty; S \rightarrow \infty$$

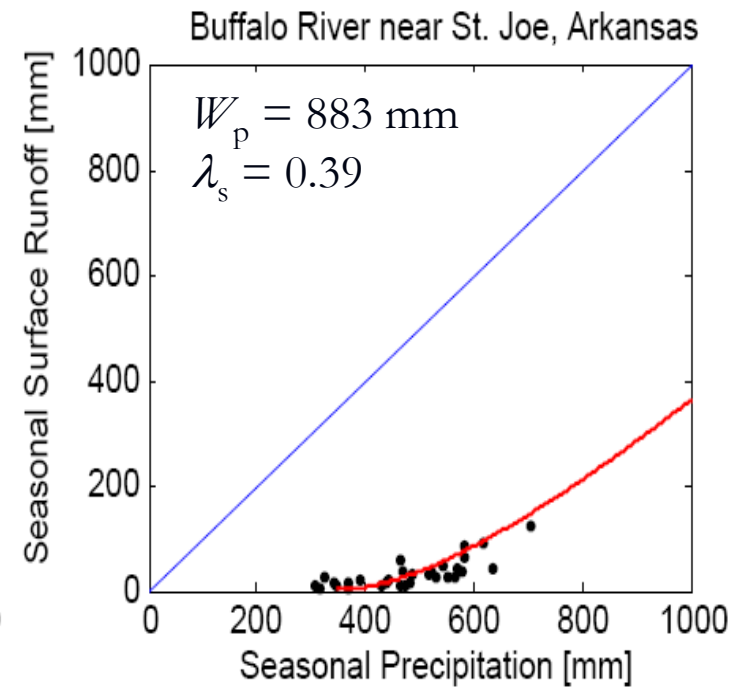
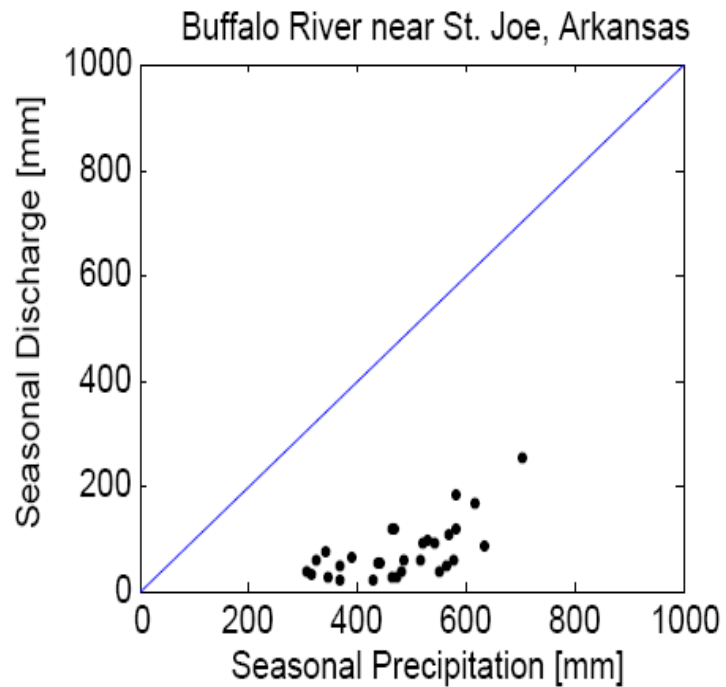
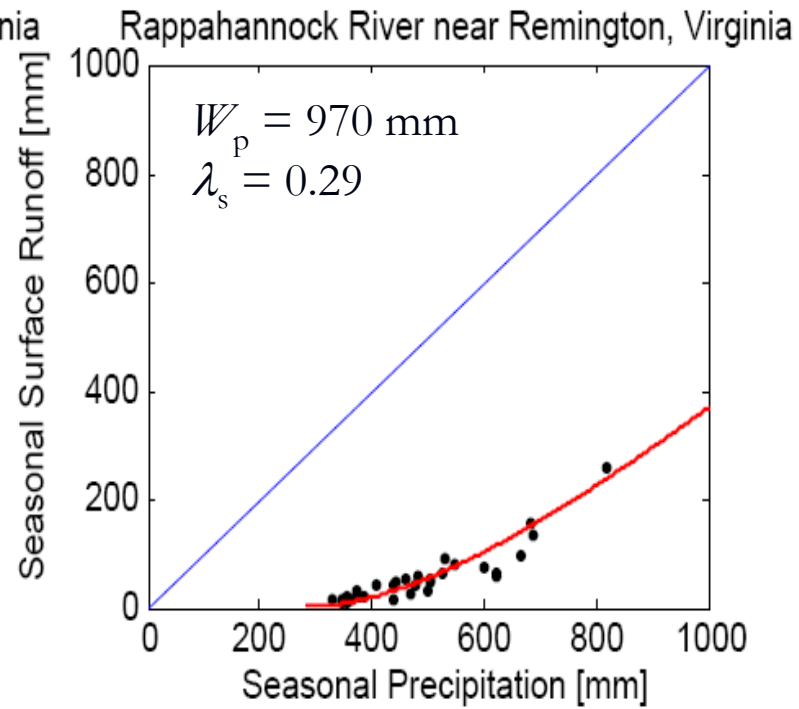
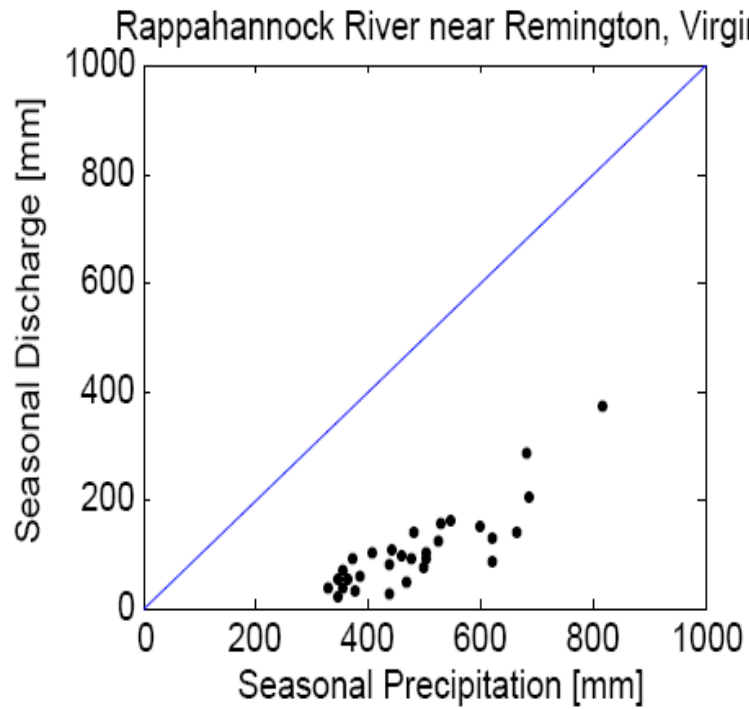
$$S = \frac{(P - \lambda_s W_p)^2}{P + (1 - 2\lambda_s)W_p}$$

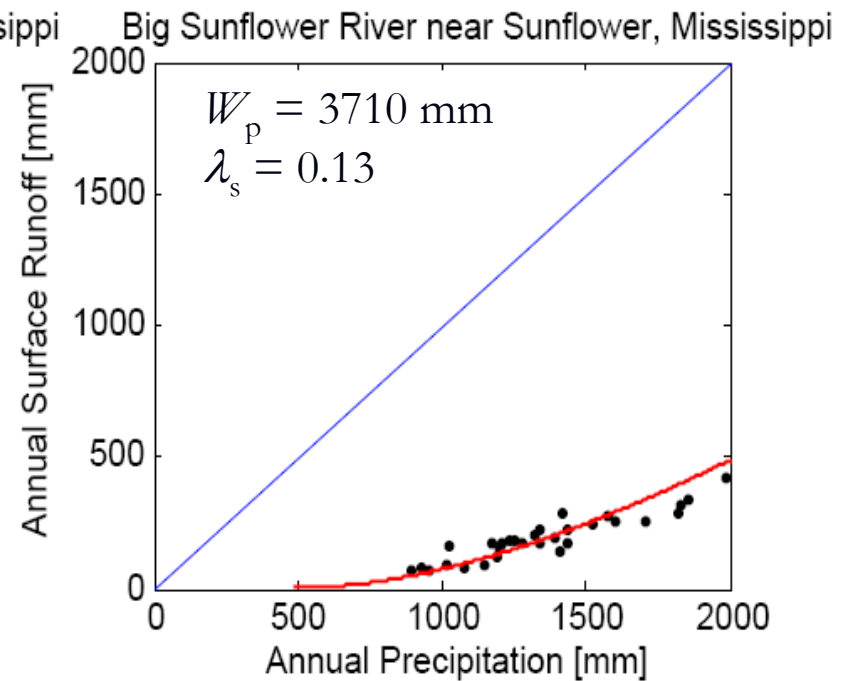
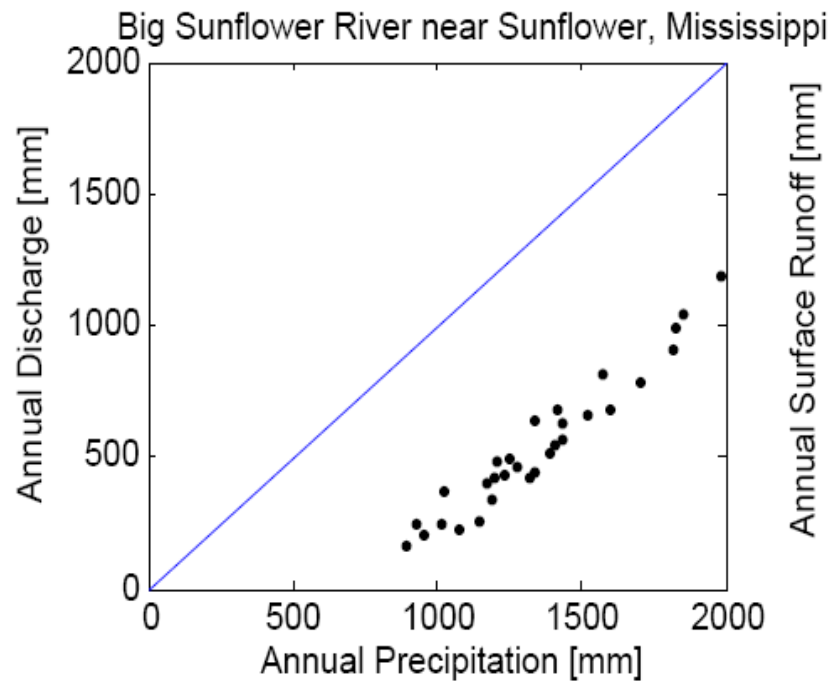
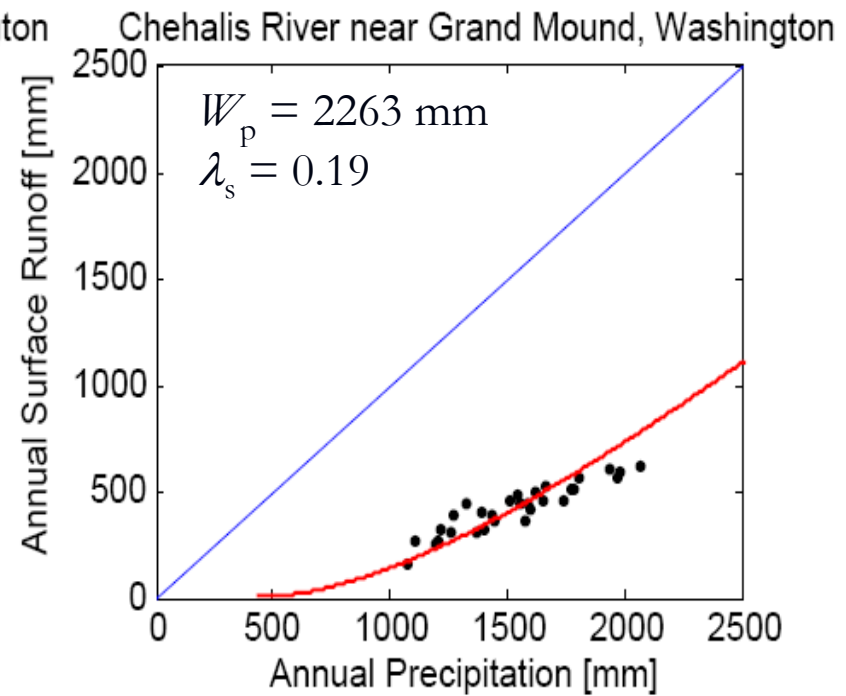
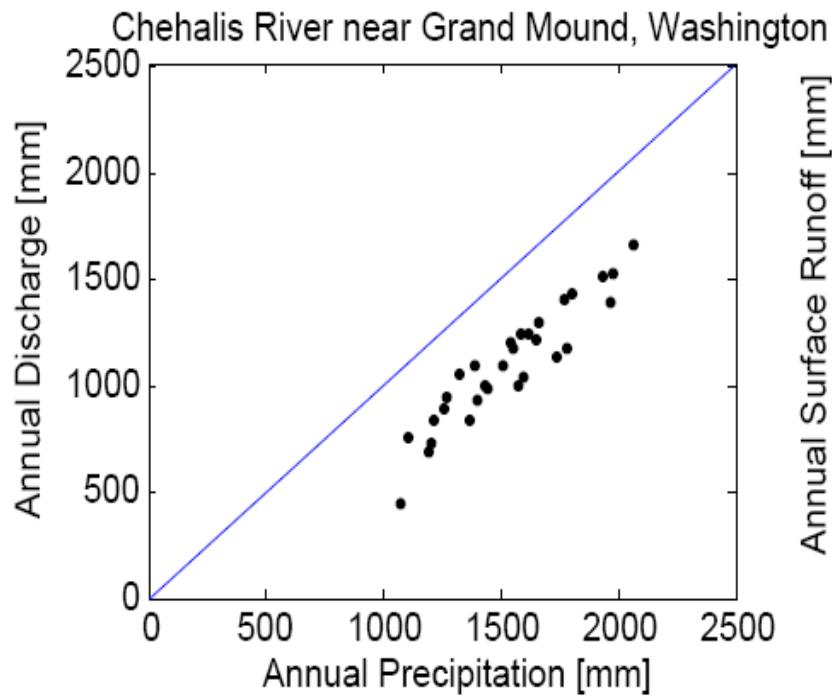
W_p : Wetting Potential (annual precipitation that can be retained by the catchment)

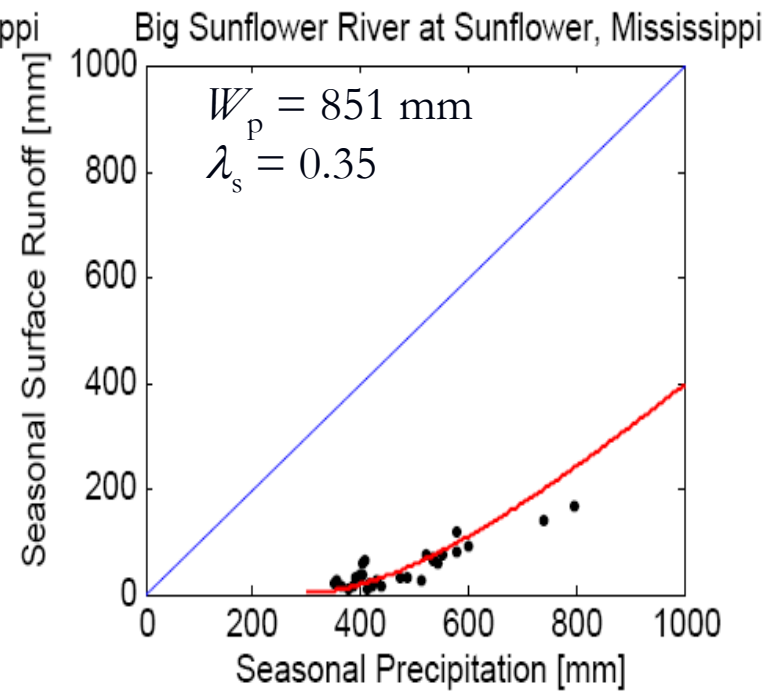
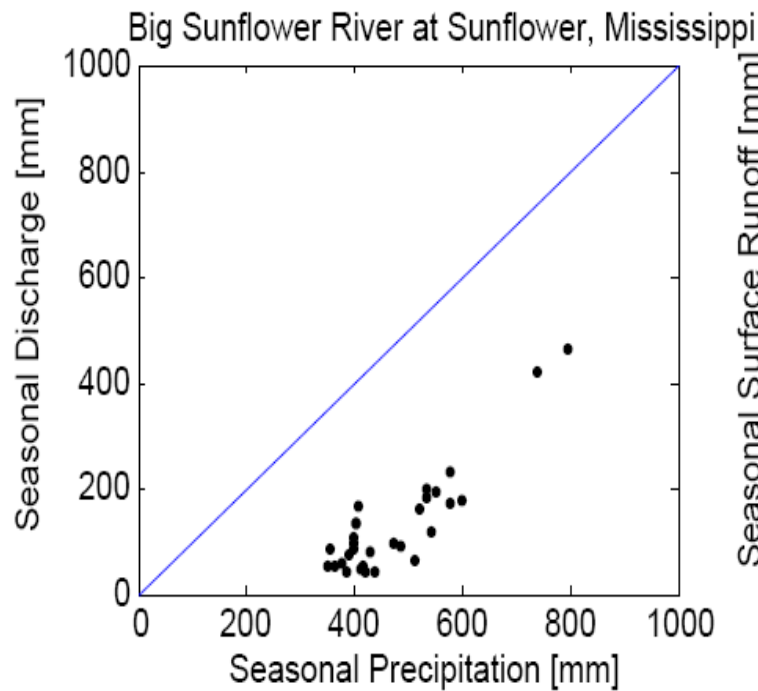
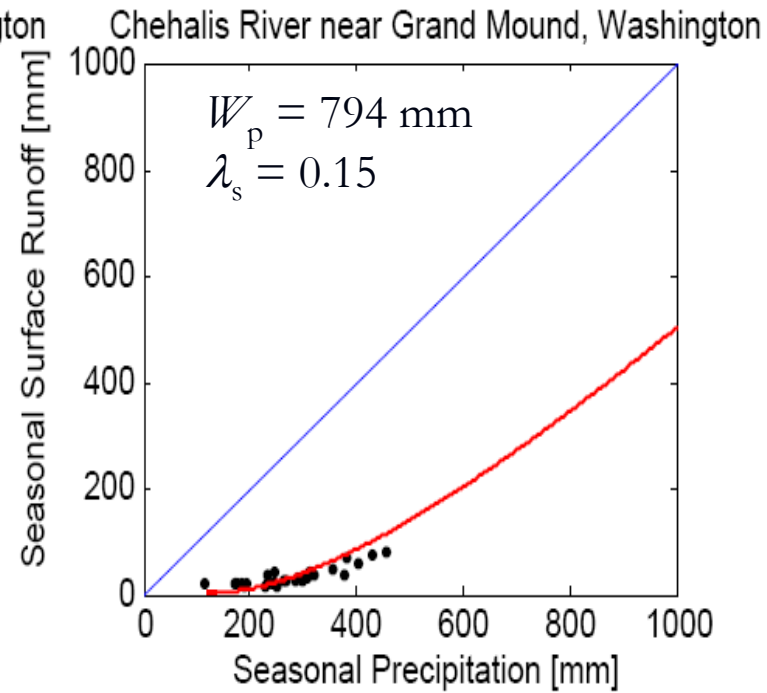
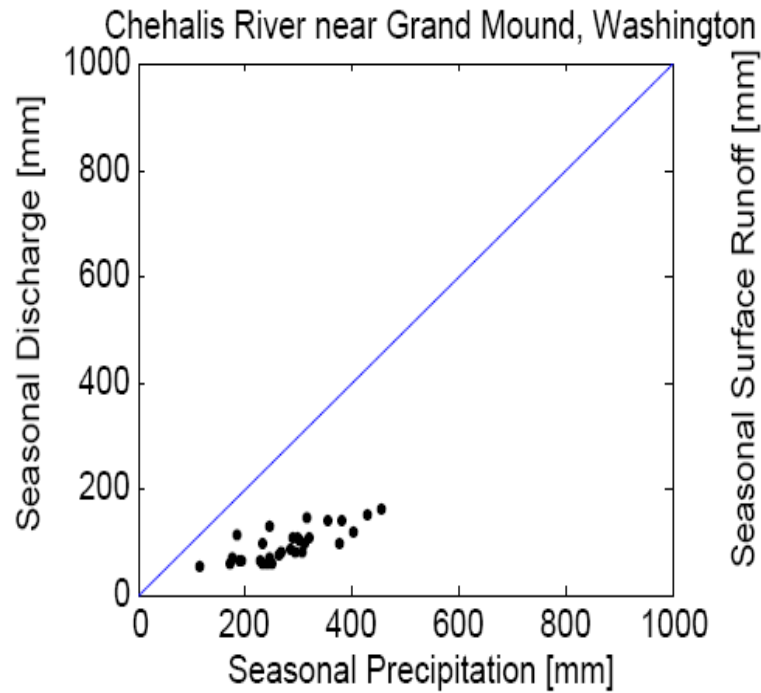
λ_s : Surface Runoff Abstraction Coefficient

Ponce and Shetty, 1995 (JoH)

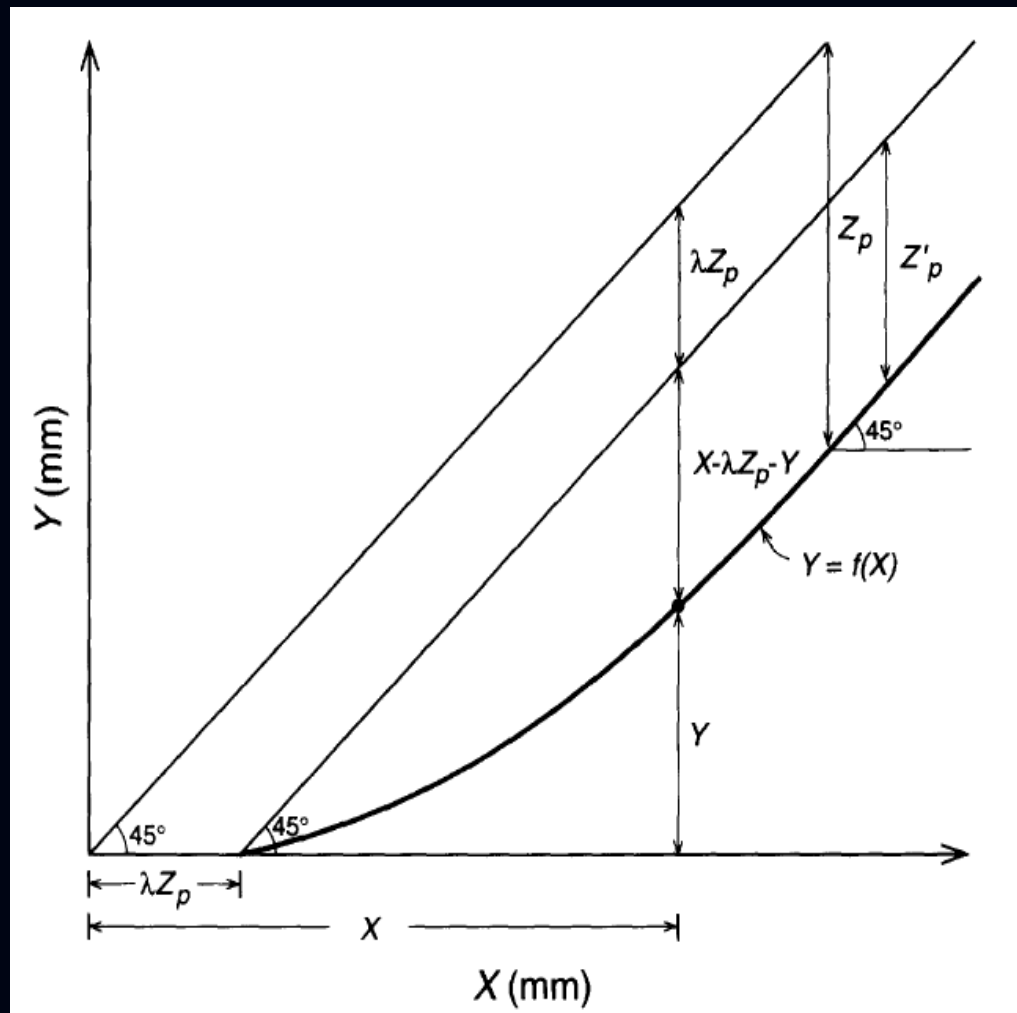








Proportionality Relations



$$X = Y + Z$$

$$Z = X - Y$$

$$Z \rightarrow Z_p \Leftrightarrow X \rightarrow \infty; Y \rightarrow \infty$$

$$W = U + V$$

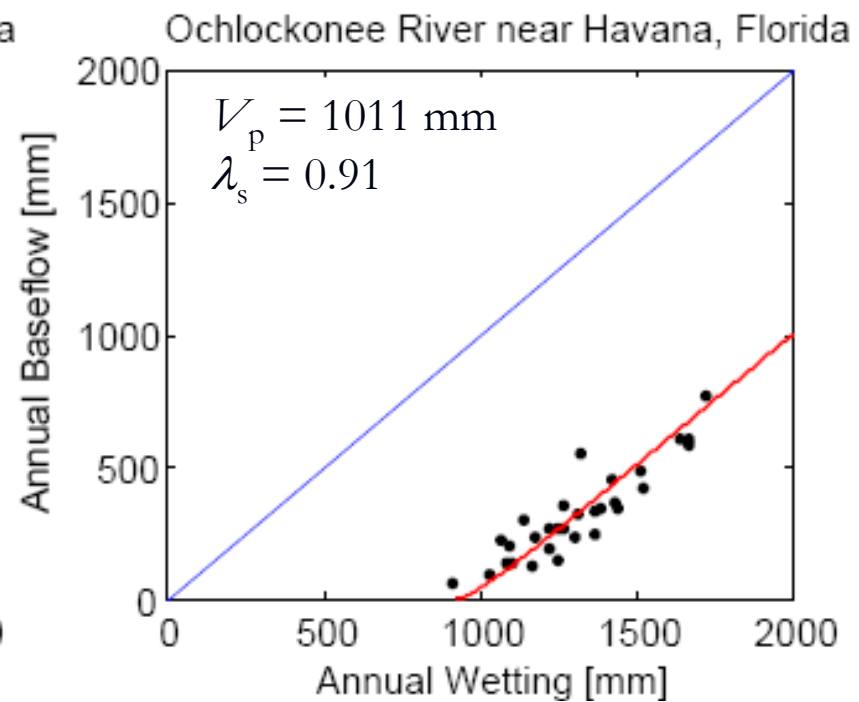
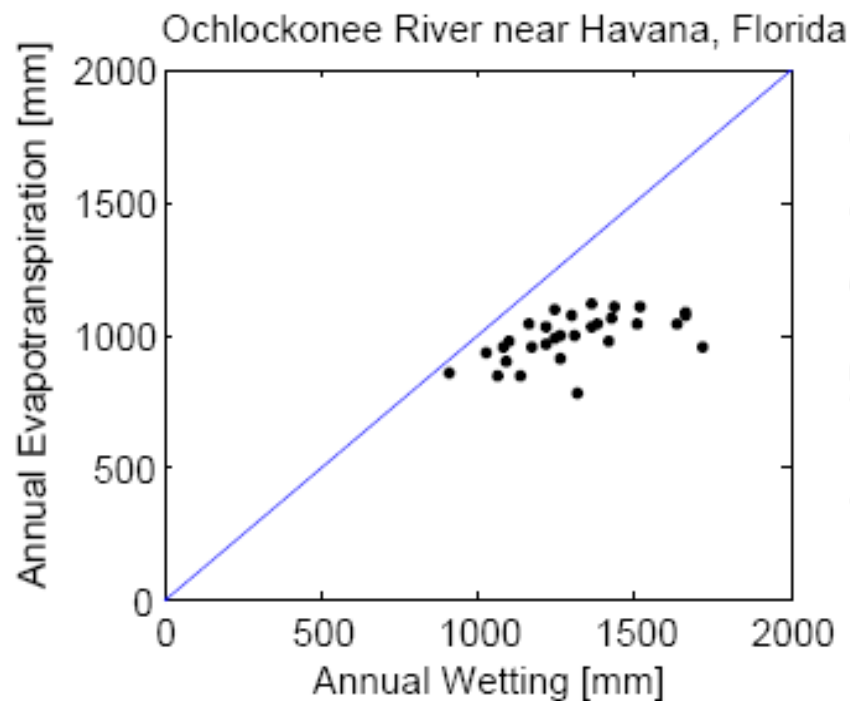
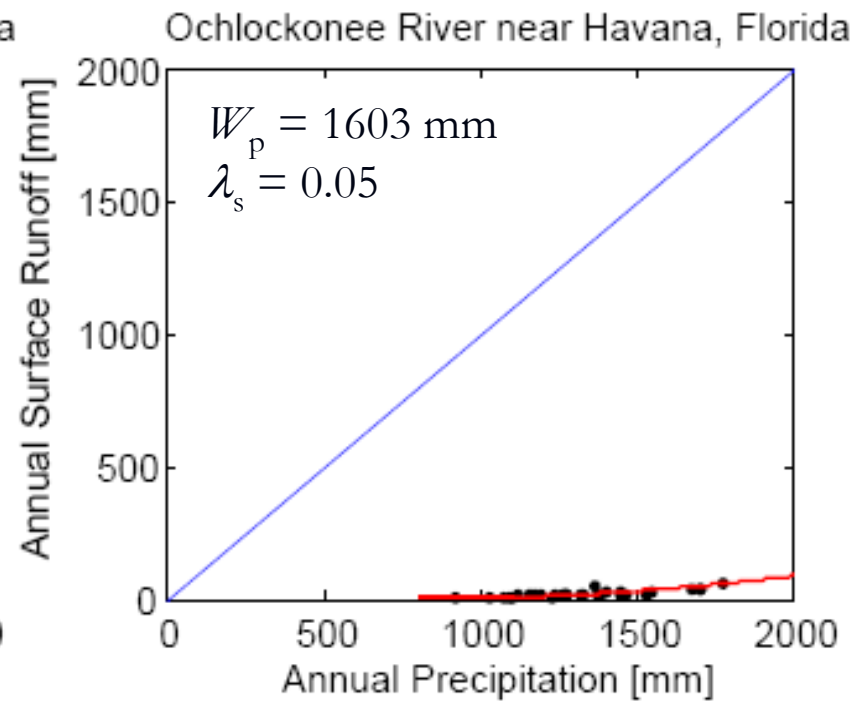
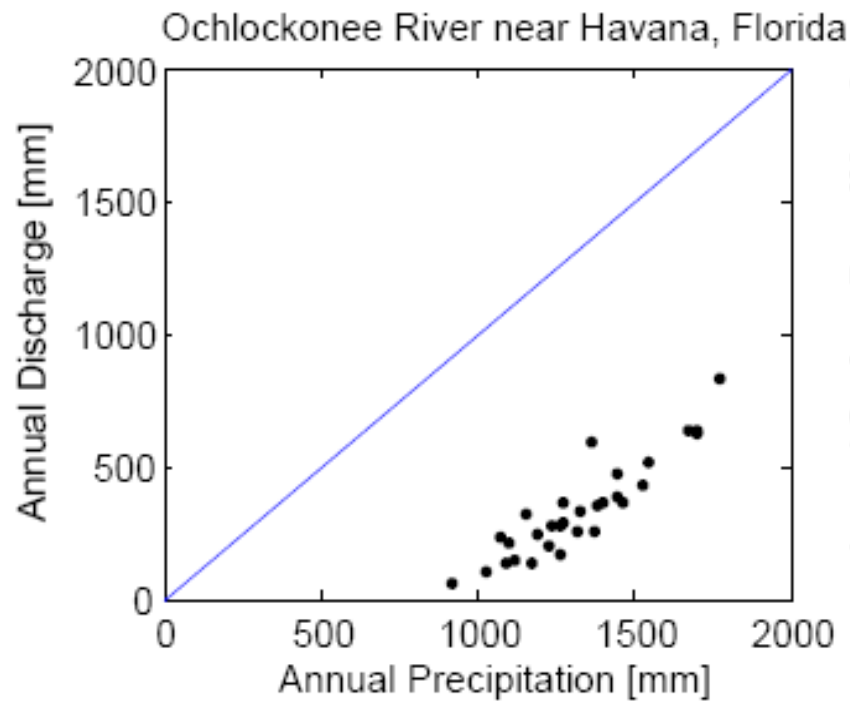
$$V = W - U$$

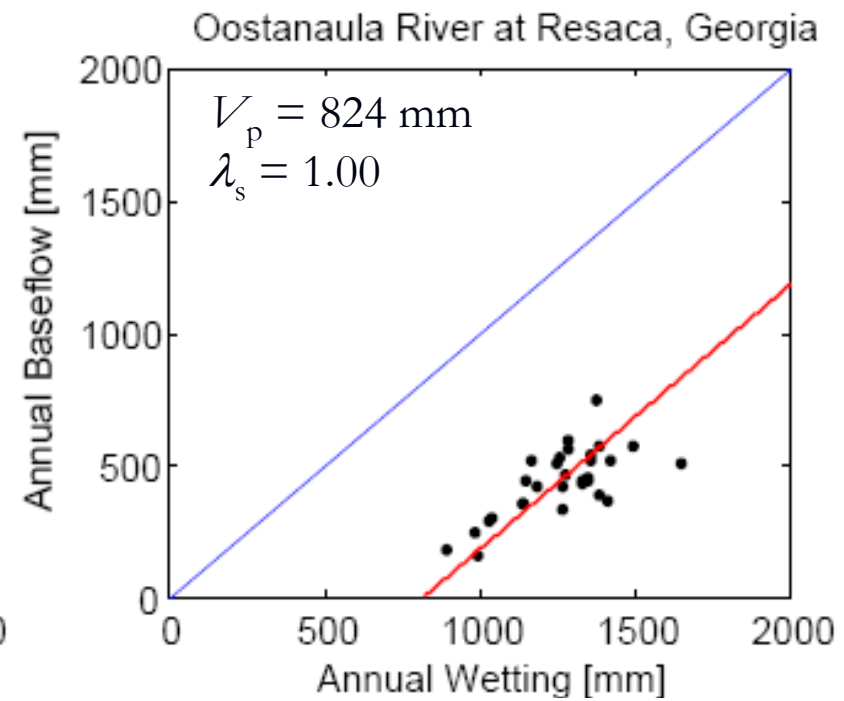
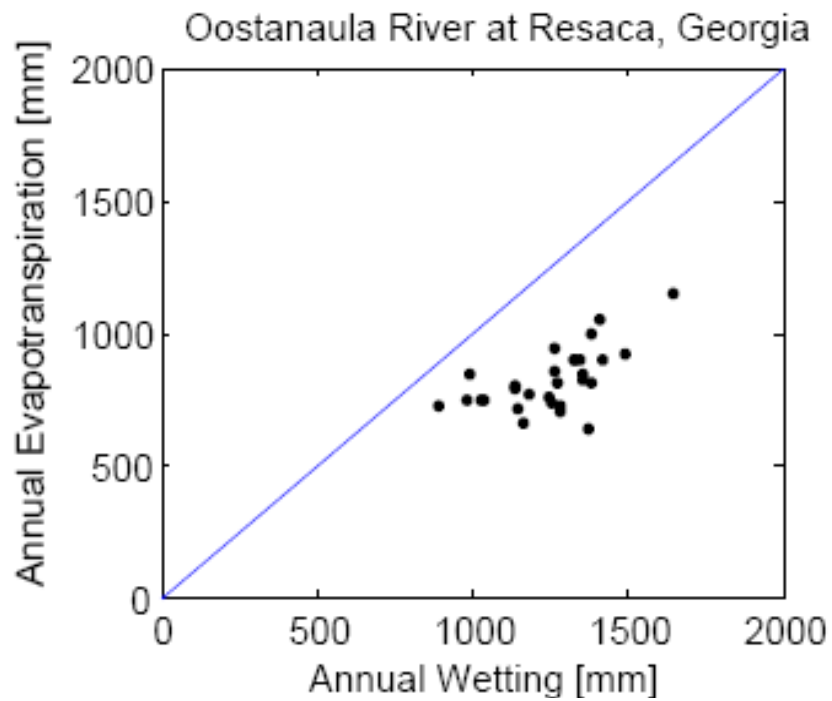
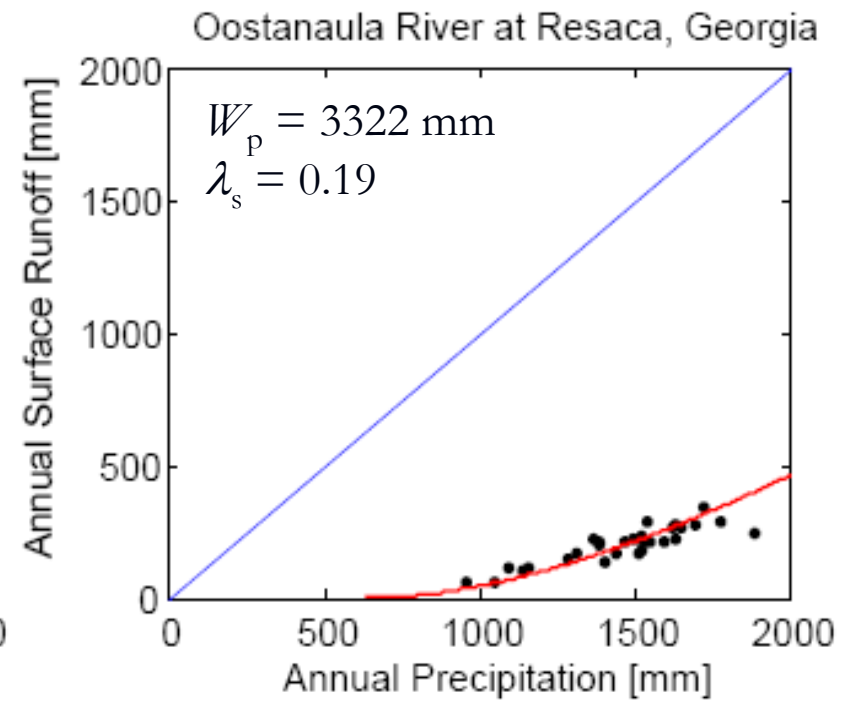
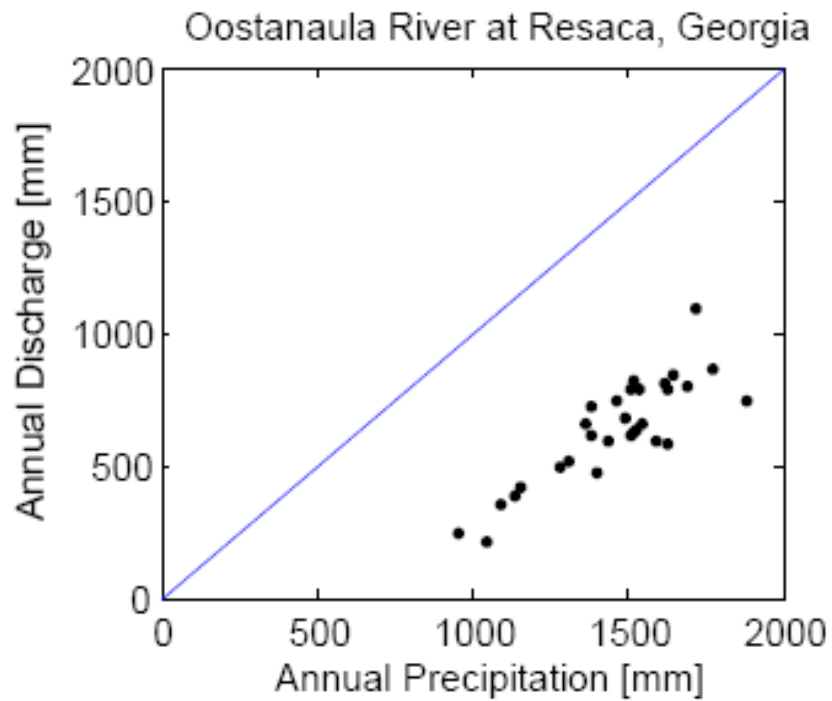
$$V \rightarrow V_p \Leftrightarrow W \rightarrow \infty; U \rightarrow \infty$$

$$U = \frac{(W - \lambda_u V_p)^2}{W + (1 - 2\lambda_u)V_p}$$

V_p : Vaporization Potential (annual wetting that can be evaporated)

λ_s : Baseflow Abstraction Coefficient



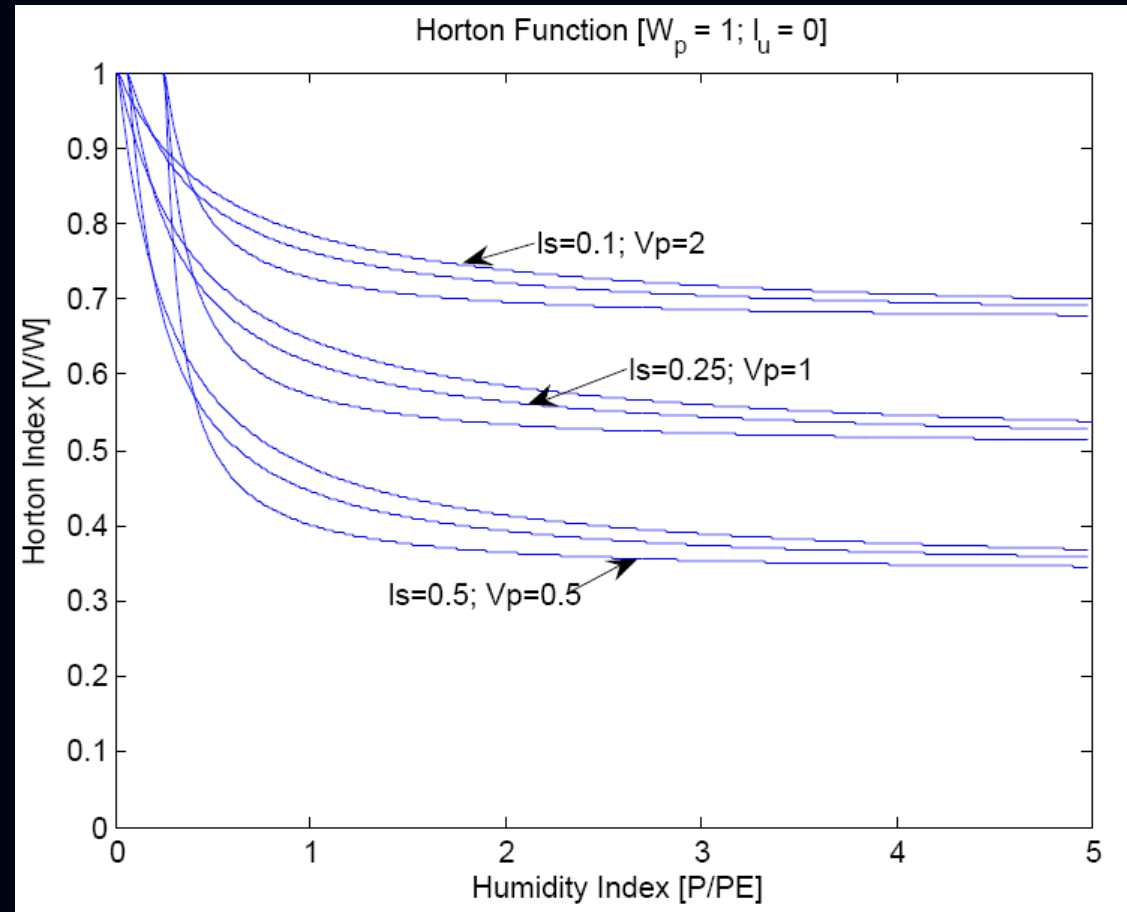


Back to the Horton Index

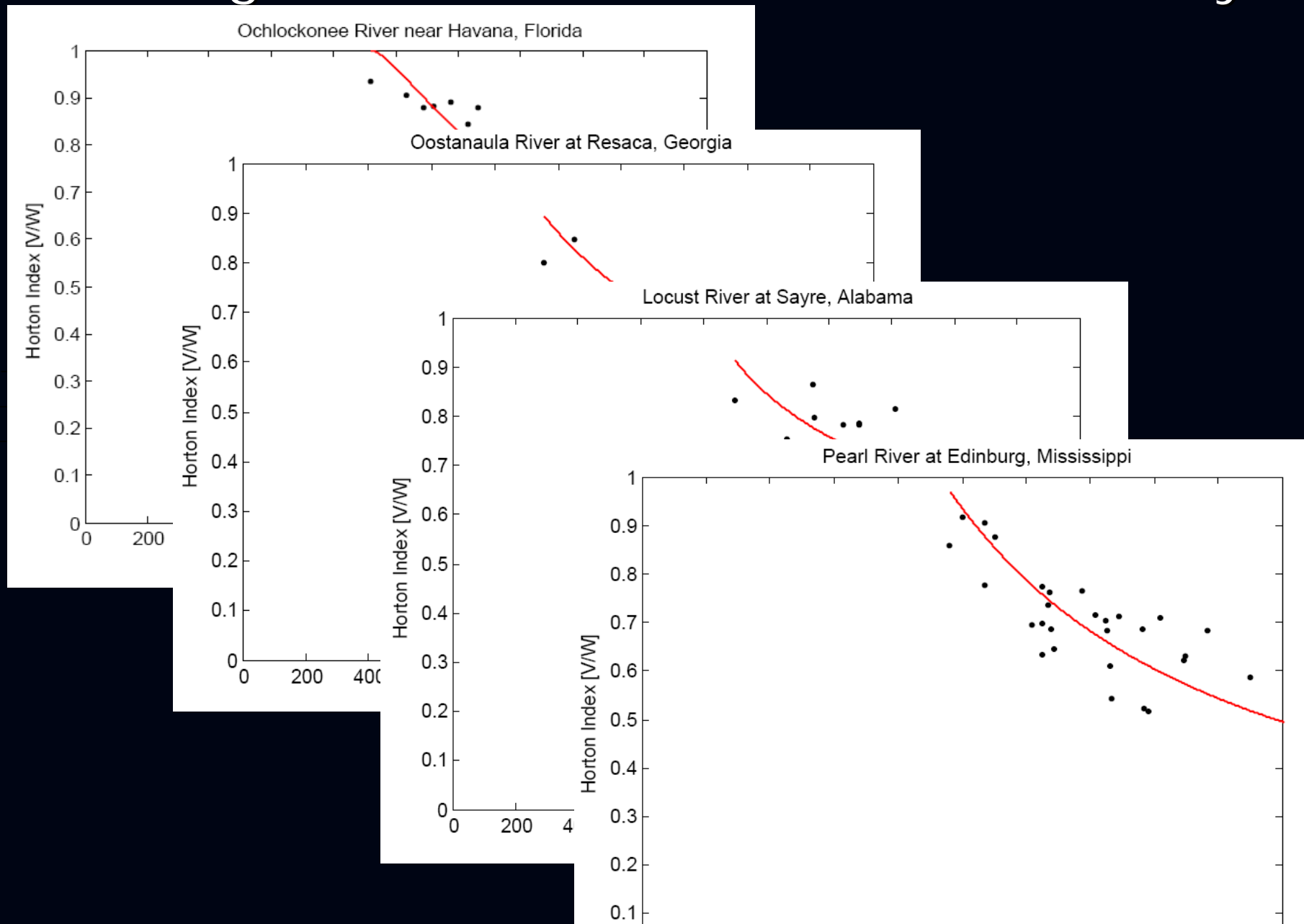
$$H = \frac{V}{W} = \frac{W - \frac{(W - \lambda_u V_p)^2}{W + (1 - 2\lambda_u)V_p}}{P - \frac{(P - \lambda_s W_p)^2}{P + (1 - 2\lambda_s)W_p}}$$

$$H = f(P, \theta)$$

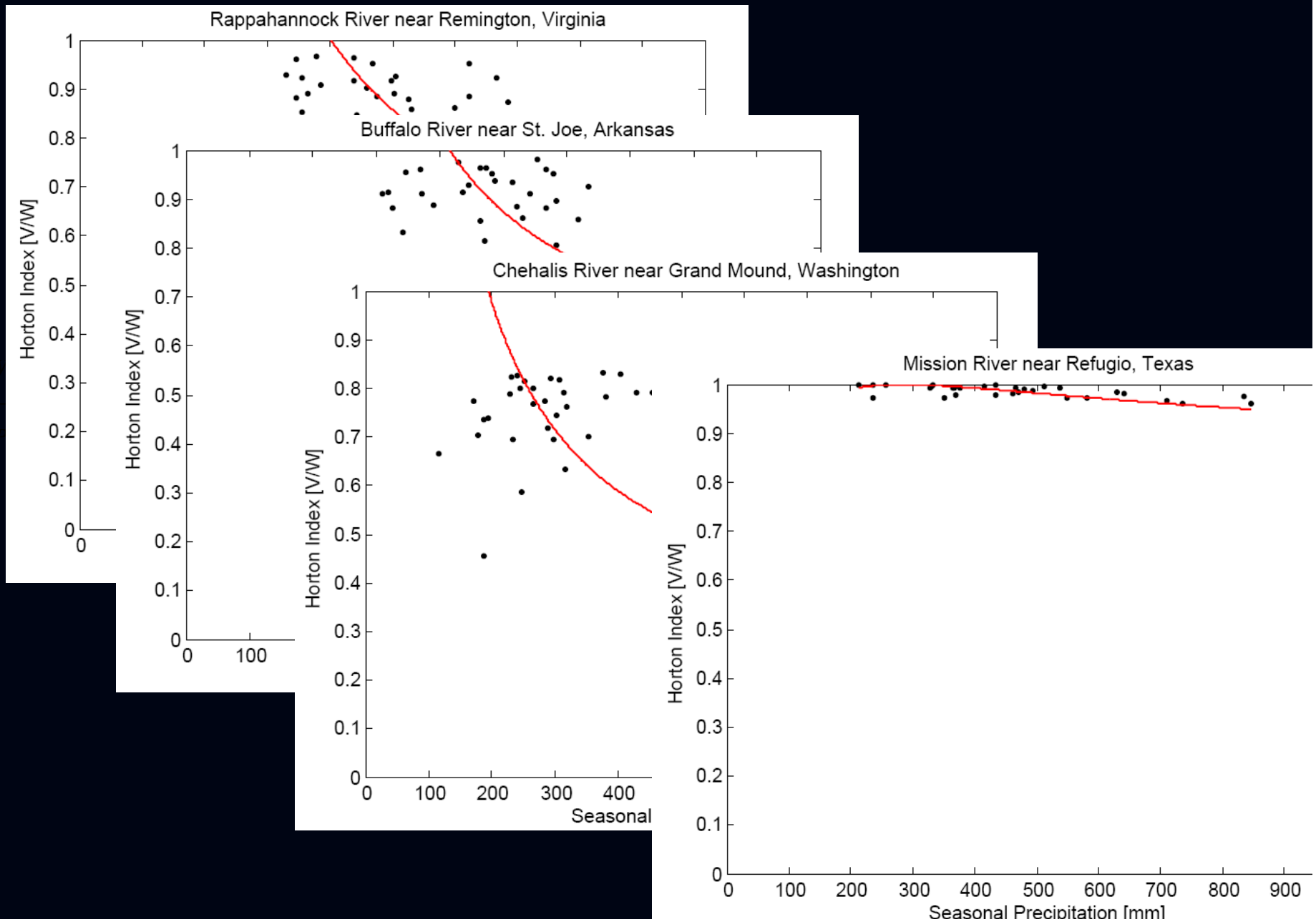
$$\theta = [V_p, W_p, \lambda_s, \lambda_u]$$



Predicting Horton Index interannual variability



Predicting Horton Index constancy



Conclusions (1)

- In semi-arid climates, the Horton index is very constant and close to 1 over the growing season, indicating that the biome WUE is constant and near maximum;
- In humid climate, the Horton index is fairly constant and its value below 1 depends on the available energy; the biome WUE depends on other factors, such as nutrients and radiation;

Conclusions (2)

- When evaluated at annual time scales, the Horton index seems to converge to a common value, similar to those observed in semi-arid climates;
- This seems to indicate that the catchment WUE converges to a common maximum WUE, in line with previous observations at the biome level;

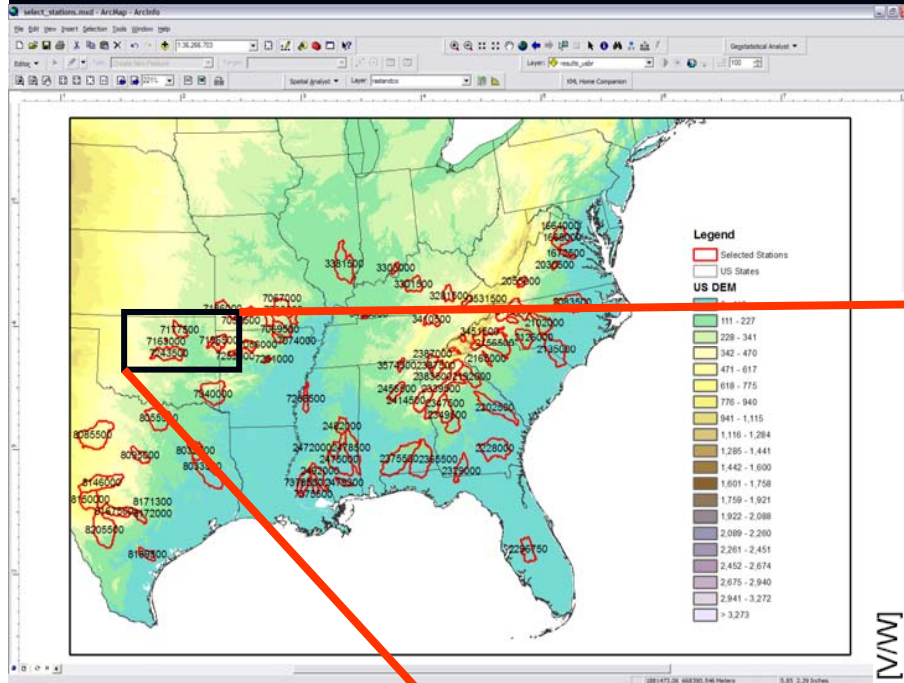
Conclusions (3)

- The interannual variability of the Horton index can be accurately reproduced using the proportionality relations of L'vovich;
- The parameters of the model indicate the catchment functioning in terms of competition between quick runoff and wetting, and between evapotranspiration and baseflow.

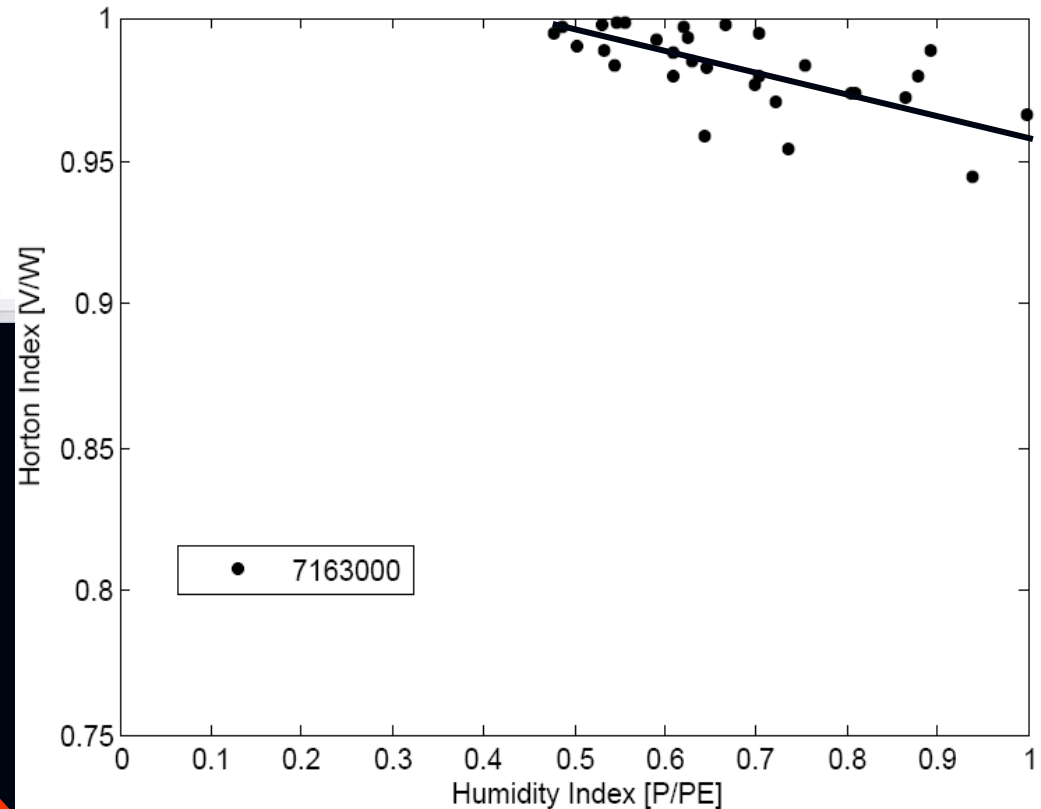
Questions?



Interannual Variability of Horton Index



Horton Index Interannual Variability
Savannahs



Ecological controls to interannual variability

